

1 New Algorithm for Rain Cell Identification and 2 Tracking in Rainfall Event Analysis

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8 **Abstract:** This study proposes a new algorithm termed rain cell identification and tracking (RCIT)
9 to identify and track rain cells from high resolution weather radar data. Previous algorithms have
10 limitations when tracking non-consequent rain cells owing to their use of maximum correlation
11 coefficient methods and their lack of an alternative way to handle the variation stages of rain cells
12 during their life cycles. To address these deficiencies, various methods are implemented in the new
13 algorithm. These include the particle image velocimetry (PIV) method for motion estimation and
14 the rain cell matching rule to obtain the stage changes of rain cells. High resolution (5-min and 1-
15 km) radar reflectivity data from three rainy days over the German federal state North Rhine
16 Westphalia (NRW) are used to evaluate the proposed algorithm. The performance of the new
17 algorithm is compared with a radar reflectivity map and verified by two object-oriented methods:
18 structure–amplitude–location (SAL) and geometric index. The verification results suggest that the
19 performance of the new algorithm is good. Application of the RCIT algorithm to the selected cases
20 shows that the inner structure of rainfall events in the experimental region present extreme value
21 distributions, with most rainfall events having a short duration with less intensity. The new
22 algorithm can effectively capture the stage changes of rain cells during their life cycles. The
23 proposed algorithm can serve as the basis for further hydro-meteorological applications such as
24 spatial and temporal analysis of rainfall events and short-term flood forecasting.

25 **Keywords:** rain cell; tracking; PIV; feature-based verification

26 1. Introduction

27 Precipitation is a key process in Earth's water circle. Acquiring explicit knowledge about its
28 inner behavior is critical to assisting us in understanding its interaction with hydrological processes.
29 Rainfall events are characterized by several elements, such as duration, intensity, velocity, and spatial
30 and temporal variability (Elena et al. 2017). The variability of rainfall events can be defined as “the
31 variability derived from having multiple spatially-distributed rainfall fields for a given point in time”
32 (Peleg et al. 2017). In hydro-meteorological applications, rainfall always varies over its life cycle; this
33 variation also differs between different types of event (e.g., convective and stratiform rainfall events).
34 As a consequence, the responses of hydrology models are sensitive to this variation. Modeling rainfall
35 events and analyzing their spatial and temporal information is necessary.

36 For rainfall event monitoring, intensity and cumulative value are the two most common indexes
37 and they are usually measured using a rain gage, which is the standard instrument for providing
38 direct observations. Nevertheless, a rain gage cannot directly detect variability in rainfall events, it is
39 also subject to errors owing to topography and wind effects. As a possible alternative, weather radar
40 has played a major role in recent years owing to its high spatial and temporal resolution. This is
41 advantageous in terms of (i) acquiring spatial and temporal patterns when modeling rainfall events
42 and (ii) undertaking short-term rainfall forecasting at fine scales. Identifying and tracking rainfall
43 events is a common task in radar-based meteorological and hydrological applications (Moseley et al.
44 2013; Novo et al. 2014; Guinard et al. 2015; Yeung et al. 2015).

45 Broadly speaking, the corresponding algorithms for radar-based rainfall event identification and
46 tracking can be classified into pixel- and object-based approaches (Zahraei et al. 2013). The pixel-
47 based approaches are also referred to as advection field tracking. These are pattern matching

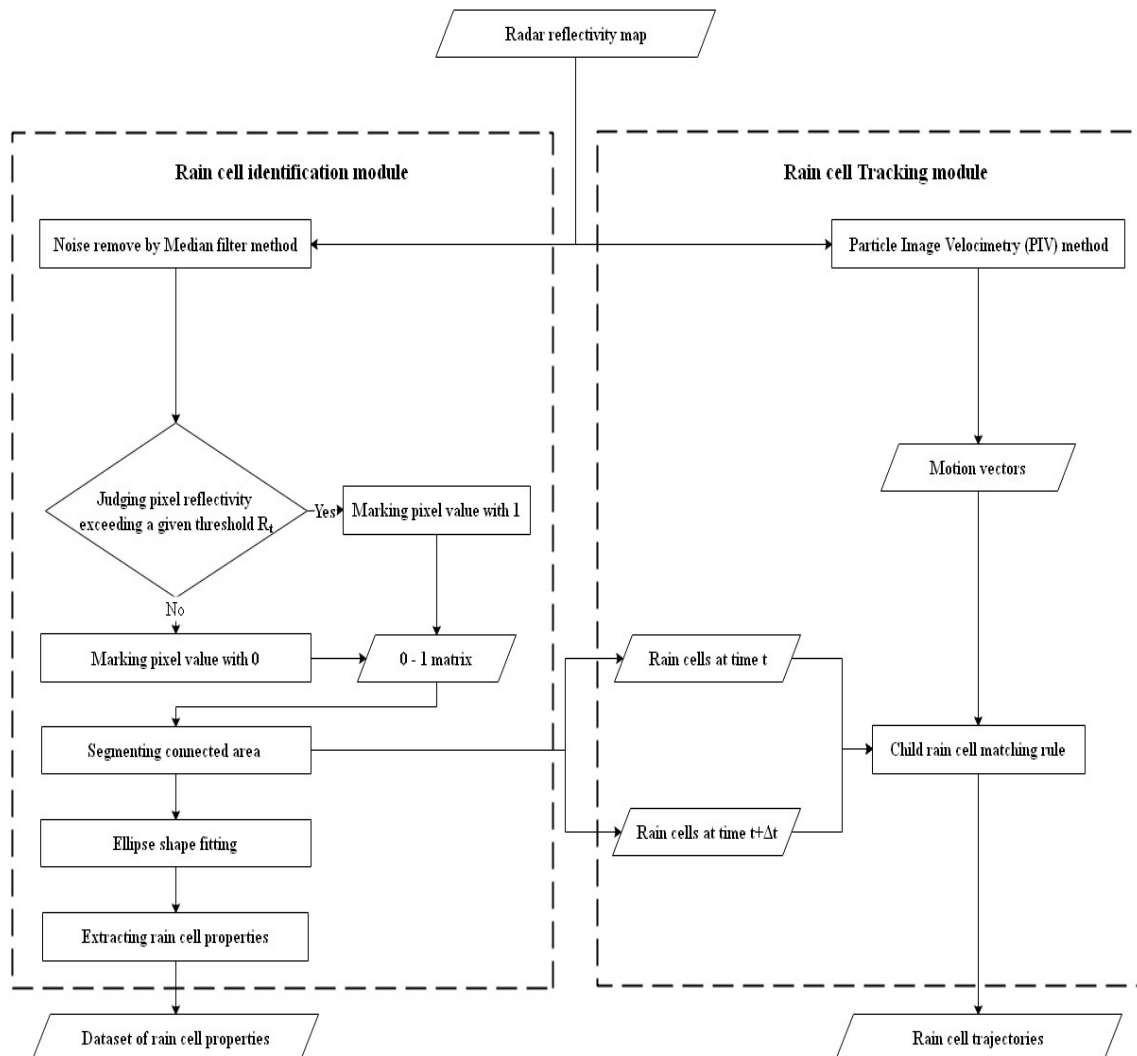
48 approaches for extracting motion vectors by searching for the maximum correlation coefficient of
49 rain cells in two consecutive radar images. Pixel-based algorithms are highly effective when they are
50 applied in convective rainfall nowcasting, which is usually found in frontal systems (Anna et al.
51 2018). Algorithms of this kind are capable of distinguishing between large scale convective and
52 stratiform rainfall, although they are not so effective for individual RCIT. Herein, a rain cell is defined
53 as the closed contours over which rainfall intensity exceeds a given threshold during one rainfall
54 event (Féral et al. 2006). Representative algorithms can be listed as follows: TREC (Rinenart and
55 Garvey 1978), TITAN (Dixon and Wiener 1993), COTREC (Li et al.1995), and SWIRLS (Li et al. 2014).
56 Object-based approaches (also termed cell tracking) include (i) a detecting algorithm for searching a
57 discrete rain cell's properties (e.g., centroid, area, echo-tops, and vertical-integrated) in consecutive
58 radar images and (ii) a matching algorithm for tracking cell motion and shape changes (e.g., merging
59 and splitting). The advantage of object-based approaches is in reflecting the dynamic of convective
60 rainfall; they are suitable for convective rain storm analysis but not effective in stratiform rainfall
61 identification. Representative algorithms can be listed as follows: SCOUT (Einfalt et al. 1990), SCIT
62 (Johnson et al. 1998), Trace3D (Handwerker 2002), and PERSiann-ForeCAST (Zahrani et al. 2013).

63 Despite the thorough application of these rain event identification and tracking algorithms, they
64 have the following deficiencies. For pixel-based approaches, motion estimates are mostly based on
65 the maximum correlation coefficient, which may yield non-continuous results when fast decay of
66 rainfall occurs. For object-based approaches, motion estimates obtained from the rain cell center of
67 mass may lack accuracy owing to the random center of mass displacement problem (Han et al., 2009).
68 This problem occurs when the shape of rain cells changes rapidly between successive radar images.
69 As a consequence of these motion estimate inaccuracies, these algorithms may also encounter
70 difficulties when handling merging and splitting scenarios (Muñoz et al., 2018), whereas a rain cell
71 can begin its life cycle by simply emerging at a location with no rain. It is also possible that a rain cell
72 can become separated from a large single rain cell or that several smaller rain cells can merge into
73 one (Moseley et al. 2013). For such reasons, a new rain cell identification and tracking algorithm
74 (RCIT) is proposed in this work. The algorithm is developed by combining the advantages of pixel-
75 and object-based approaches and is able to handle the problem of detecting the stage changes of rain
76 cells.

77 This paper is organized as follows. An introduction to the study area and radar data is given in
78 Section 2. The RCIT algorithm is illustrated in Section 3. Section 4 presents structure, amplitude, and
79 location (SAL) and geometric verification results and practical applications of the algorithm to North
80 Rhine Westphalia (NRW) rainfall events. In Section 5, the main conclusions and further expectations
81 of this work are given.

82 2. Rain Cell Identification and Tracking Algorithm - RCIT

83 The aim of the RCIT is to analyze rainfall events by fully utilizing the merits of weather radar.
84 The inputs to the proposed algorithm are radar reflectivity maps and the outputs are the properties
85 of rain cells such as area, intensity, and trajectory. The RCIT involves two modules: rain cell
86 identification and tracking, as presented as in Figure 1.



87

88

Figure 1. Step illustration of rain cell identification and tracking algorithm.

89

2.1. Rain Cell Identification Module

90 Similar to the object-based algorithms, the rain cell identification module of the RCIT is based
 91 on discerning a connected domain above a given threshold. As presented in the left portion of Figure
 92 1, each selected radar reflectivity map in the Cartesian coordinates is initially filtered by the median
 93 filtering method (Anoraganingrum 1999) to remove noisy pixels (pixels with abnormally high
 94 reflectivity). Then, pixels above a given reflectivity threshold are assigned the value one, with the
 95 remainder assigned the value zero. A segmenting process is implemented to assemble and cluster
 96 pixels sharing the same reflectivity threshold into a connected area. In the segmenting process, the
 97 following rules suggested by Peleg and Morin (2012) are obeyed: (i) If the reflectivity of a rainy pixel
 98 is lower than a given threshold R_t , then it is set to null. (ii) For each rainy pixel and its eight neighbors,
 99 if more than five of them are null, then it is set to null. (iii) If the pixel is spurred, then it is set to null.
 100 Herein, spur pixels are those isolated pixels whose reflectivity is different to others along the
 101 horizontal and vertical directions in the labeled binary image. (iv) If the area of a connected region is
 102 smaller than 9 km², then it is ignored. All the segmented regions are then labeled and fitted with an
 103 ellipse shape. Their properties are extracted and stored in a relational database. The extracted
 104 properties are as follows:

105

i) Area [km²] - Sum value for the number of pixels contained in one rain cell.

106

ii) Areal rainfall depth [mm] - Cumulative precipitation of one rain cell over a 5-minute interval.

- 107 iii) Maximum intensity [mm.h⁻¹] - Peak intensity of one rain cell.
 108 iv) Areal mean rainfall depth [mm.km²] - Ratio of the areal rainfall depth and area.
 109 v) Eccentricity - Ratio of minor and major axes, which are acquired from the fitted ellipse. Used
 110 to describe the shape of one rain cell with a value range from 0 to 1.
 111 vi) Center of mass [km] - Center of mass of a rain cell, which is weighted by the reflectivity of rainy
 112 pixels.
 113 Property calculation: Areal rainfall depth, maximum intensity, and areal mean rainfall depth are
 114 based on the reflectivity (Z) and rain rate (R) converting function: $Z = aR^b$.

115 2.2. Rain Cell Tracking Module

116 The rain cell tracking module is established based on a hybrid approach, as illustrated in the
 117 right portion of Figure 1. In the first procedure, motion vectors are estimated by implementing the
 118 particle image velocimetry (PIV). This is an optical method of flow visualization that is used to obtain
 119 instantaneous velocity measurements and related properties of fluids (Merzkirch 2001; Adrian 2005;
 120 Westerweel et al. 2013). It consists of a class of flow measuring mechanisms that are characterized by
 121 recording the displacement of small particles embedded in a region of fluid. Figure 2 shows a PIV
 122 application in motion vector estimation. In the first step, a window box of size $r \times r$ is initially defined,
 123 which divides radar images into several sub-blocks. In the second step, a searching distance, $d =$
 124 $2 \times v_{\max} + 1$, is defined, where v_{\max} is the preset maximum velocity. The minimum quadric difference
 125 (MQD), as suggested by Gui and Merzkirch (1996), is employed in searching the optimal grid points
 126 at time $t + \Delta t$, as in Equation (1):

$$127 \quad MQD(\Delta x, \Delta y) = \sum_{i=1}^N \sum_{j=1}^N |R_1(X_i, Y_j) - R_2(X_i + \Delta x, Y_j + \Delta y)| \quad (1)$$

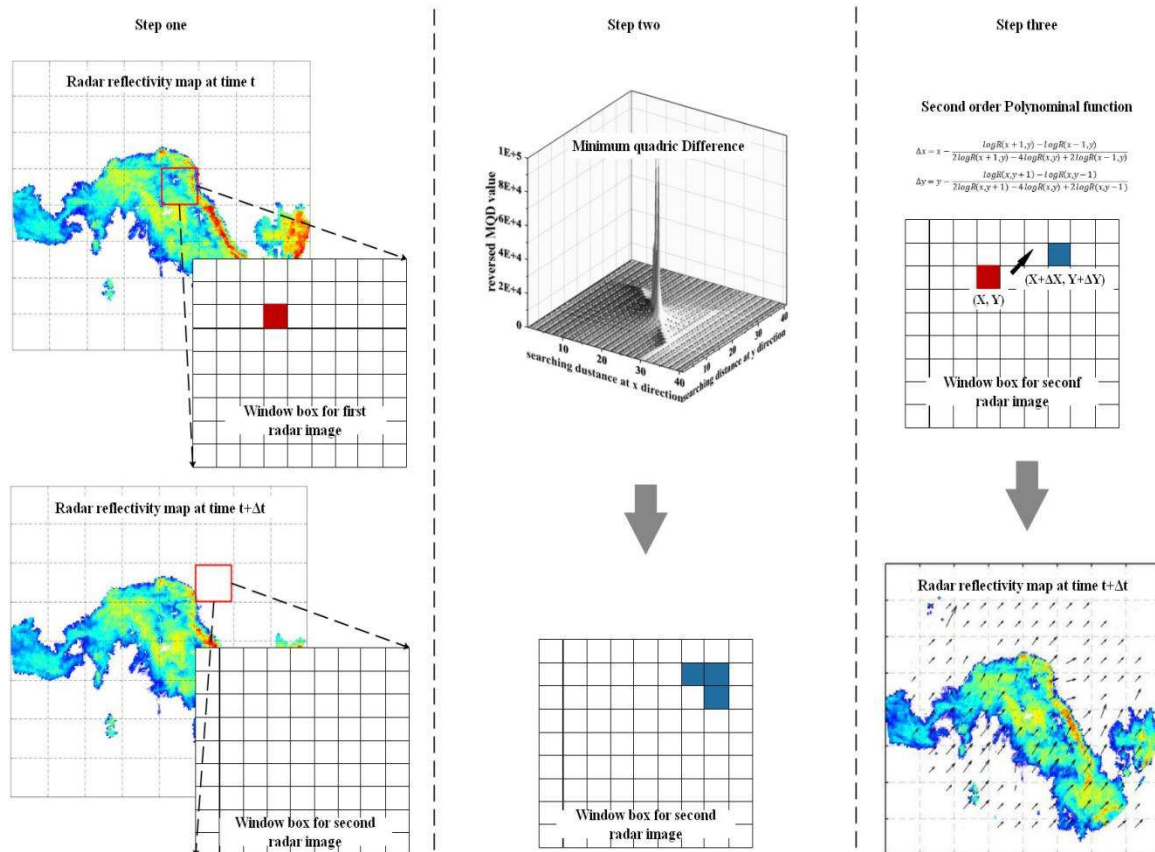
128 where $R_1(X_i, Y_j)$ and $R_2(X_i, Y_j)$ are the reflectivity of grid points contained within the window boxes of
 129 radar images at time t and $t + \Delta t$, respectively; Δx and Δy ($\Delta x, \Delta y \in d$) are the locations of minimum
 130 reflectivity difference in the horizontal and vertical directions separately. The minimum reflectivity
 131 differences of grid points within the window boxes are reversed to simplify the calculation. In order
 132 to guarantee that the solitary peak locations can be calculated, Δx and Δy are corrected separately in
 133 the horizontal and vertical directions by fitting a second-order polynomial to the logarithm of the
 134 maximum reflectivity of the grid point and its three direct neighbors, as in Equation (2). In this way,
 135 the optimal grid points at time $t + \Delta t$ are identified, with their locations presented as $(x + \Delta x - \frac{d+1}{2}, y +$
 136 $\Delta y - \frac{d+1}{2})$. In the final step, the calculated motion vectors are smoothed by the median filter algorithm.

$$137 \quad \Delta x = x - \frac{\log R(x+1, y) - \log R(x-1, y)}{2 \log R(x+1, y) - 4 \log R(x, y) + 2 \log R(x-1, y)} \quad (2a)$$

$$138 \quad \Delta y = y - \frac{\log R(x, y+1) - \log R(x, y-1)}{2 \log R(x, y+1) - 4 \log R(x, y) + 2 \log R(x, y-1)} \quad (2b)$$

139 In the second procedure, rain cells at time t and $t + \Delta t$ are identified. Finally, a child rain cell
 140 matching rule is applied for identifying the most-matched rain cells. The child rain cell matching
 141 scheme considers the stage changes of rain cells between successive radar images (e.g., merge, split,
 142 growth, and decay), using certain indexes for determination such as overlap, area diversification,
 143 distance, and angle difference of center of mass. Before introduction of the rain cluster matching rule,
 144 certain definitions were identified: (i) for two radar images at time t and $t + \Delta t$, rain cells identified
 145 from radar images at t are termed parent cells and (ii) rain cells identified from radar images at $t + \Delta t$
 146 are termed child cells. These definitions can be depicted as follows:

- 147 (1) A boundary box of a parent cell is defined, with a horizontal length of $[10+\max(\text{pos}_x),$
 148 $\min(\text{pos}_x)-10]$ and vertical length of $[10+\max(\text{pos}_y), \min(\text{pos}_y)-10]$, where pos_x and pos_y are
 149 Cartesian coordinates of pixels in the parent cell.
- 150 (2) The number of child cells falling into the boundary box is determined and their
 151 properties, e.g., area, areal rainfall depth, max intensity, areal mean rainfall depth, and
 152 center of mass, are selected.
- 153 (3) If only one child cell is searched in the boundary box and it overlaps with a parent
 154 cell, then it is the most-matched rain cell. If this child cell does not overlap with a parent cell
 155 and the distance and angle difference for the center of mass between it and the parent cell
 156 are less than $3 \times \text{mean}(V_{\text{motion_vector}})$ and $3 \times \theta_{\text{motion_vector}}$, it is also the most-matched rain cell,
 157 where $\text{mean}(V_{\text{motion_vector}})$ and $\theta_{\text{motion_vector}}$ are the mean value of velocity and the prevailing
 158 direction of the motion vector, respectively.
- 159 (4) If two or more child cells fall into the boundary box without overlapping a parent cell,
 160 the matching rule is changeless; however one extra condition is included, i.e., child cells
 161 whose areas have minimum absolute differences with the parent cell are the most-matched
 162 rain cells.



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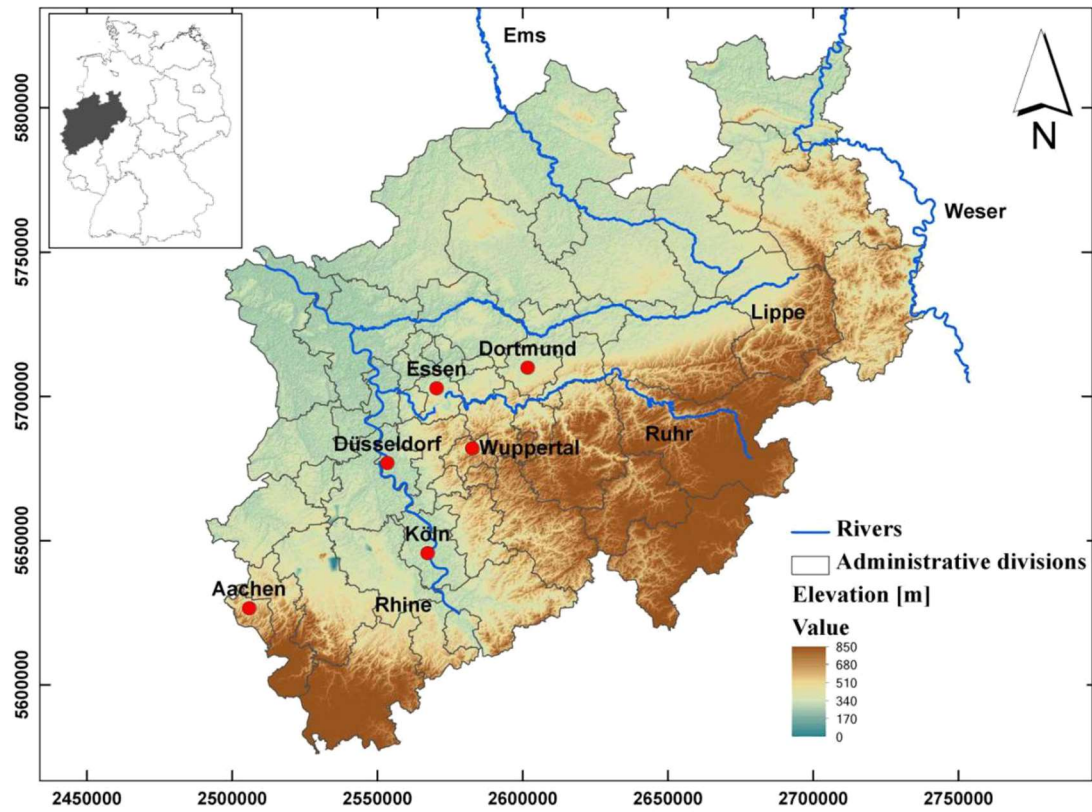
164 **Figure 2.** Illustration of PIV application in rain cell motion estimation. Step one: window boxes with
 165 the area of $r \times r$ are defined (rectangular with red color); step two: for any grid point in the window
 166 box at a previous timepoint (red block), the MQD algorithm is applied to deduce the minimum
 167 reflectivity differences, and grid points with minimum value in the next window box are identified
 168 (blue blocks); step three: the solitary locations of reversed MQD value, Δx and Δy , are corrected by
 169 applying Equation (2). The locations of the optimal grid point at the next timepoint are calculated
 170 using the second polynomial function, and the global motion vectors are extracted and smoothed by
 171 the median filter method.

172 3. Study Area and Data

173 The study area is the federal state of NRW (Figure 3). It is bordered by the German federal states
174 of Lower Saxony to the north and north-east, Hessen to the east, and Rhineland–Palatinate to the
175 south, and by the countries of Belgium to the south-west and the Netherlands to the west. NRW
176 includes upland regions of North Eifel in the south and mountains of the Sauerland in the south-east.
177 There are five rivers in this study area: Rhine, Ruhr, Ems, Lippe, and Weser. Two main types of
178 landscape can be found in NRW, namely the North German lowlands, with elevations just a few
179 meters above sea level, and the North German low mountain range, with elevations of up to 850 m.
180 The lowland areas comprise the Rhine–Ruhr area, which is one of the largest metropolitan areas, with
181 a population of approximately 10 million. The circulation pattern of NRW is mainly affected by the
182 air mass from the Atlantic Ocean along the direction toward the mainland. When arriving at the
183 southern high mountain regions, the air mass stops and rises; this leads to more cloudiness and
184 precipitation. On the eastern side of the mountain regions, drier air masses result in less cloudiness
185 and less precipitation.

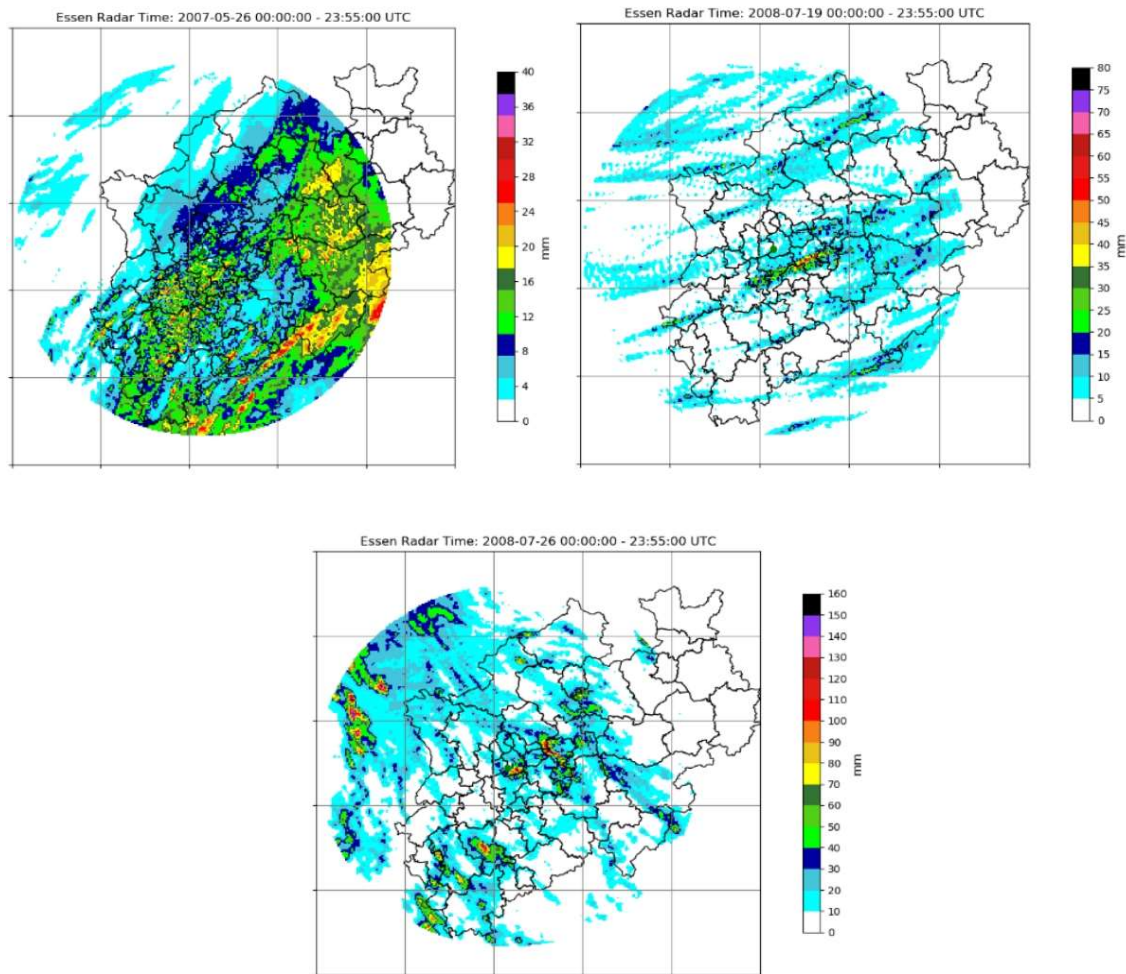
186 Radar data were obtained from the Essen radar deployed in Essen City, NRW. The Essen radar
187 has been deployed over the study area and is a part of the radar network of the German weather
188 service (DWD). The Essen radar is a dual-polarimetric C-band Doppler radar. It delivers radar
189 volume scans (frequency: 800/1200 Hz, maximum range: 124 km) every 15 minutes for the Doppler
190 velocity, together with intensity volume scans (frequency: 500 Hz, maximum range: 256 km) and
191 precipitation scans (frequency: 600 Hz, maximum range: 150 km) every 5 minutes for the
192 precipitation echo, with a high spatial resolution of 1 km in range and 1° in azimuth. For this study,
193 the Essen radar provided precipitation scans with an elevation of 0.8° and a range of 128 km. The
194 output reflectivity was selected with a plan position indicator (PPI) display type.

195 As radar measures precipitation in an indirect manner, the quality of radar data must be
196 carefully checked. The sources that affect the quality of the Essen radar data include ground clutter
197 and speckle, beam blockage, and attenuation. The corresponding quality correction methods in this
198 work follow Golz et al. (2006). After quality checking, an open source package “Wradlib”
199 (Heistermann et al. 2013) was applied to project the raw radar image onto a 256 × 256 km² Cartesian
200 map with 1-km resolution. In total, 864 radar reflectivity images for three rainy days (May 26, 2007;
201 July 19, 2008; and July 26, 2008), including some recorded convective rainfall events, were selected to
202 evaluate the proposed algorithm. The daily rainfall distributions of these rainy days are shown in
203 Figure 4.



204
205
206

Figure 3. Study area of North Rhine Westphalia (NRW) and its location in Germany. The main administrative cities are marked with red dots.



207

208 **Figure 4.** Daily rainfall distribution for selected rainy days. Radar reflectivity was converted into rain
 209 rate according to DWD standard Z–R relationship.

210 4. Results Analysis and Discussion

211 In the algorithm application process, the reflectivity threshold for rain cell identification was set
 212 to 19 dBZ, which was based on the classification of DWD presented by Weusthoff and Hauf (2008),
 213 as in Table 1. For the German weather radar system, two common Z–R relationships were used by
 214 Weusthoff and Hauf (2008): one was categorized for the RADOLAN product and the other uses
 215 constant a and b with values of 256 and 1.42, respectively. Although the DWD has stated that the
 216 categorized relationship statistically shows better results over long time periods, the standard
 217 relationship can be more compatible with local cases when a correction factor is added (Einfalt and
 218 Frerk 2012). Based on the above considerations, we applied the DWD standard Z–R relationship to
 219 radar reflectivity–rain rate conversion in the application cases.

220 **Table 1.** Conversion of radar reflectivity to rain rates using Z–R relationship ($a = 256$, $b = 1.42$) with
 221 thresholds according to the classification of the DWD.

Z [dBZ]	R [mm.h ⁻¹]	Rain Rate [mm.5 min ⁻¹]
> 55	> 150	> 12.5
46–55	35–150	2.92–12.5
37–46	8.1–35	0.68–2.92
28–37	1.9–8.1	0.16–0.68
19–28	0.4–1.9	0.03–0.16
7–19	0.06–0.4	0.005–0.03

222 4.1. Performance Assessment of RCIT Algorithm

223 Grid-point related error measurement is problematic for the rain cell tracking algorithm. A
 224 classic example illustrating this problem is the well-known “Double Penalty” problem, in which
 225 prediction of a precipitation object at the correct size and structure might yield very poor verification
 226 scores. For example, one rain cell is displaced slightly in space but the categorical verification scores
 227 heavily penalize such a situation. In traditional verification methods, a displacement simply leads to
 228 a false alarm, and it is also very poorly rated owing to its large root mean squared error (Davis et al.
 229 2006). On the other hand, despite a great deal of effort in the statistical validation of grid-based
 230 rainfall estimated results, verification associated with the geometry patterns of rain cells has not been
 231 well researched or applied.

232 As an alternative, feature-based verification methods have been built upon the idea of
 233 identifying rainfall events as “objects”. With this perspective, simulated and observed rainfalls are
 234 not compared directly at the same location; rather, objects of interest are extracted from simulated
 235 and observed data and then compared together so that verification statistics are obtained. A number
 236 of spatial verification methods have been proposed (Ebert 2008; Gilleland et al. 2009). In the present
 237 work two feature-based verification methods, SAL and geometric index, are implemented to verify
 238 the performance of the RCIT algorithm. A detailed introduction to the SAL and geometric index
 239 methods is presented in Appendix A. The data set for comparison was a simulated radar reflectivity
 240 map from the RCIT algorithm (termed *sim_ref*) and an observed radar reflectivity map (termed
 241 *obs_ref*).

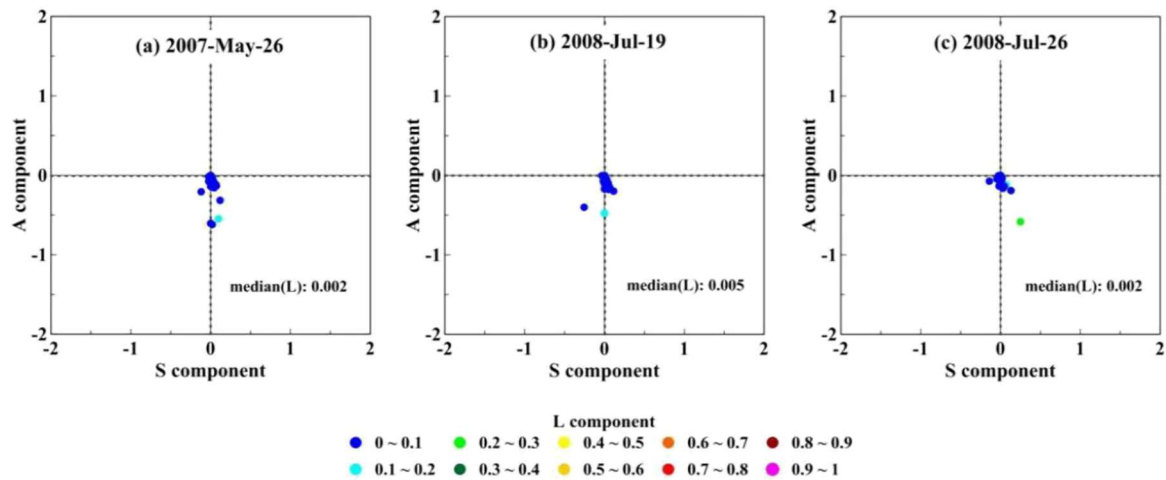
242 Figure 5 shows the SAL verification results, which are arranged based on the three selected rainy
 243 days. For each SAL plot presented in Figure 5, the vertical axis denotes the A component, the
 244 horizontal axis denoted the S component, the dots represent values of the S and A components, and
 245 the color scale of the dots denotes the L component. Median values of the SAL components are
 246 presented as dashed lines. It can be seen that all S and A values are concentrated close to zero, as are
 247 most L component values. Table 2 gives the S, A, and L index values for the geometric index objects
 248 from the *sim_ref* and *obs_ref* data sets. All values were again organized based on the three selected
 249 rainy days. It is evident that the index differences between the geometric index objects from the
 250 *sim_ref* and *obs_ref* data sets were less than 0.05.

251 **Table 2.** Results of three geometric index components of geometric index verification objects for RCIT-
 252 simulated radar reflectivity maps (*sim*) and observed radar reflectivity maps (*obs*). Values are sorted
 253 at 5, 50, and 75 percentile levels.

Selected cases		C index			S index			A index		
		25%	50%	75%	25%	50%	75%	25%	50%	75%
May 26, 2007	obs	0.934	0.957	0.977	0.22	0.325	0.509	0.102	0.198	0.417
	sim	0.966	0.979	0.992	0.27	0.378	0.579	0.135	0.271	0.53
July 19, 2008	obs	0.847	0.895	0.938	0.143	0.217	0.29	0.031	0.071	0.118
	sim	0.907	0.943	0.969	0.154	0.233	0.297	0.043	0.086	0.134
July 26, 2008	obs	0.897	0.93	0.955	0.154	0.245	0.374	0.045	0.116	0.213
	sim	0.936	0.965	0.997	0.189	0.285	0.385	0.077	0.149	0.238

254 The SAL verification results suggest that the shape of most SAL objects from the *sim_ref* data
 255 set was the same as that for SAL objects from the *obs_ref* data set (except for a few cases that were
 256 slightly large and flat). The converted rainfall volume for some SAL objects from the *sim_ref* data set
 257 was less than that from the *obs_ref* data set; the origin that the rain cell area threshold used in the
 258 RCIT algorithm was 9 km², with rain cells less than this threshold ignored. However, the converted
 259 rainfall volume of most SAL objects from the *sim_ref* data set was close to that from the *obs_ref* data
 260 set. Location differences of SAL objects between the *sim_ref* and *obs_ref* data sets were not obvious.
 261 Geometric index verification results indicated that the geometric pattern of geometric index objects
 262 from the *sim_ref* data set was approximately the same as that for objects from the *obs_ref* data set
 263 (except for connectivity). Differences in the C index for geometric index objects from the two data

264 sets were obvious, which may have been because the median filter method applied in the RCIT
 265 algorithm smoothed abnormal pixels in the radar reflectivity map. In general, the RCIT algorithm
 266 performed well based on the two feature-based verification methods.



267
 268 **Figure 5.** Value distributions of the three SAL components. Dashed lines in the vertical and horizontal
 269 directions in each sub-figure represent the median values of the S and A components, respectively,
 270 and dot color represents the of L component value. Results are sorted by selected rain days: (a) May
 271 26, 2007; (b) July 19, 2008; (c) July 26, 2008.

272 4.2. Application of RCIT Algorithm in Rainfall Event Analysis

273 There were 10,346 rain cells identified from the radar reflectivity maps. Descriptive statistics of
 274 their properties are given in Table 3. It was found that a high standard deviation existed for the areas
 275 of these rain cells, with values ranging from 9 to 18,734 km² (most were less than 38 km²). For areal
 276 rainfall depth, values ranged from 0.36 to 8861 mm and a high standard deviation again existed. For
 277 the maximum intensity property, a high standard deviation (34.08 mm.h⁻¹) was also found; the
 278 median value was 2.83 mm.h⁻¹ and the range of values was 0.48 to 397.75 mm.h⁻¹. Areal mean rainfall
 279 depth was from 0.04 to 4.4 mm.km². Eccentricity ranged from 0 to 1, with a median value over 0.5.

280 **Table 3.** Descriptive statistics of rain cell properties. Indexes used for the statistics are minimum
 281 value, maximum value, standard deviation, and median value.

Property	Statistical properties			
	Minimum value	Maximum value	Standard deviation	Median value
Area	9	18734	1391	38
Areal rainfall depth	0.36	8861	559.9	4.4
Max intensity	0.04	33.2	2.8	0.24
Areal mean rainfall depth	0.04	4.4	0.3	0.1
Eccentricity	0	0.99	0.17	0.84

283 Inner structures of the selected events were described by statistically analyzing the physical and
 284 geometric properties of the rain cells. RCIT simulation results indicated that the properties (e.g., area,
 285 areal rainfall depth, max intensity, and areal mean rainfall depth) of the identified rain cells presented
 286 a wide range of values. The shape of the rain cells was somewhat elliptical, with a median value over
 287 0.5. Histograms of the \log_{10} -transformed rain cell properties are given as in Figure 6. To determine
 288 the best theoretical distributions describing the empirical distributions, a multi-goodness of fit testing
 289 (GOF) approach combined with the Akaike information criterion (AIC), the Bayesian information
 290 criterion (BIC), and the Kolmogorov–Smirnov (K–S) test methods was applied (see Appendix B).
 291 Figure 7 shows their fitted cumulative distributions. Empirical distributions of the \log_{10} -transformed
 292 properties (area, areal rainfall depth, maximum intensity, and areal mean rainfall depth) could be
 293 fitted with the generalized Pareto distribution (GPD) presented in Equation (6), and the extreme
 294 value distribution (EVD) was found to fit the eccentricity property shown in Equation (7).

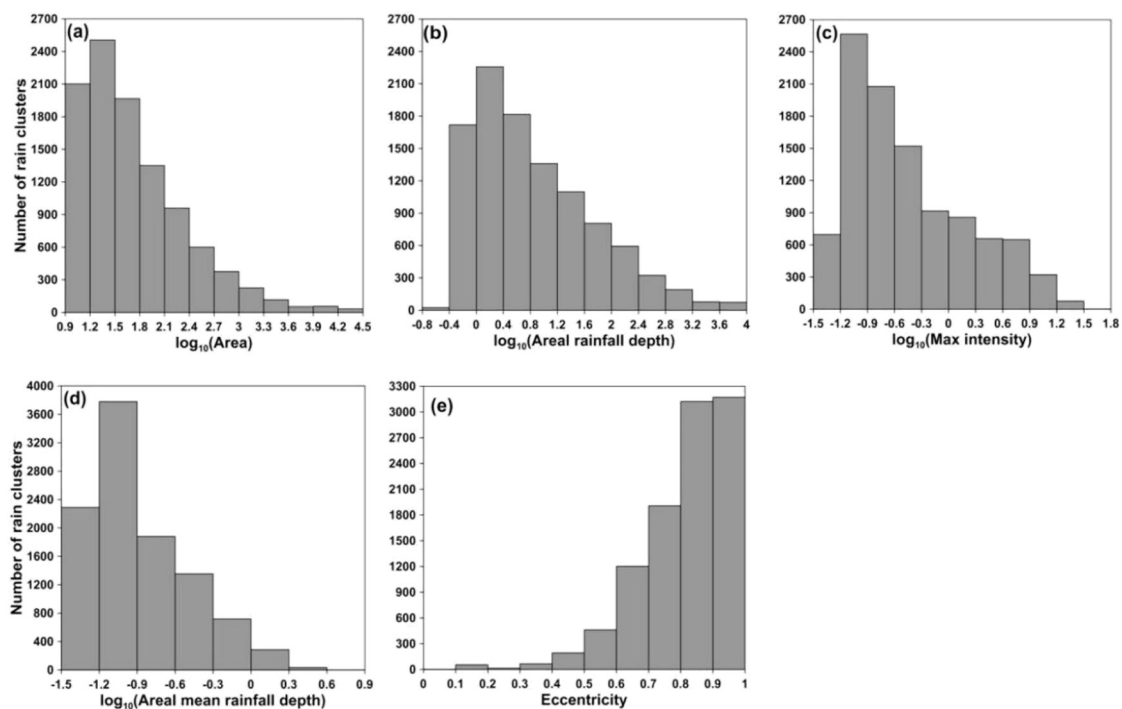
$$295 \quad f(x|k, \mu, \alpha) = \left(\frac{1}{\alpha}\right) \left(1 + k \frac{(x-\mu)}{\alpha}\right)^{-1-\frac{1}{k}} \quad (6)$$

296 where k is the shape parameter, and μ and α are location and scale parameters, respectively. For
 297 $\mu < 0$, k is above zero, and for $\mu < x < \alpha$, k is below zero. At the limit for $k = 0$, the GPD is the
 298 exponential distribution.

$$299 \quad f(x|k, \mu, \alpha) = \left(\frac{1}{\alpha}\right) e^{-\left(1+k\frac{(x-\mu)}{\alpha}\right)^{\frac{1}{k}}} \left(1+k\frac{(x-\mu)}{\alpha}\right)^{-1-\frac{1}{k}} \quad (7)$$

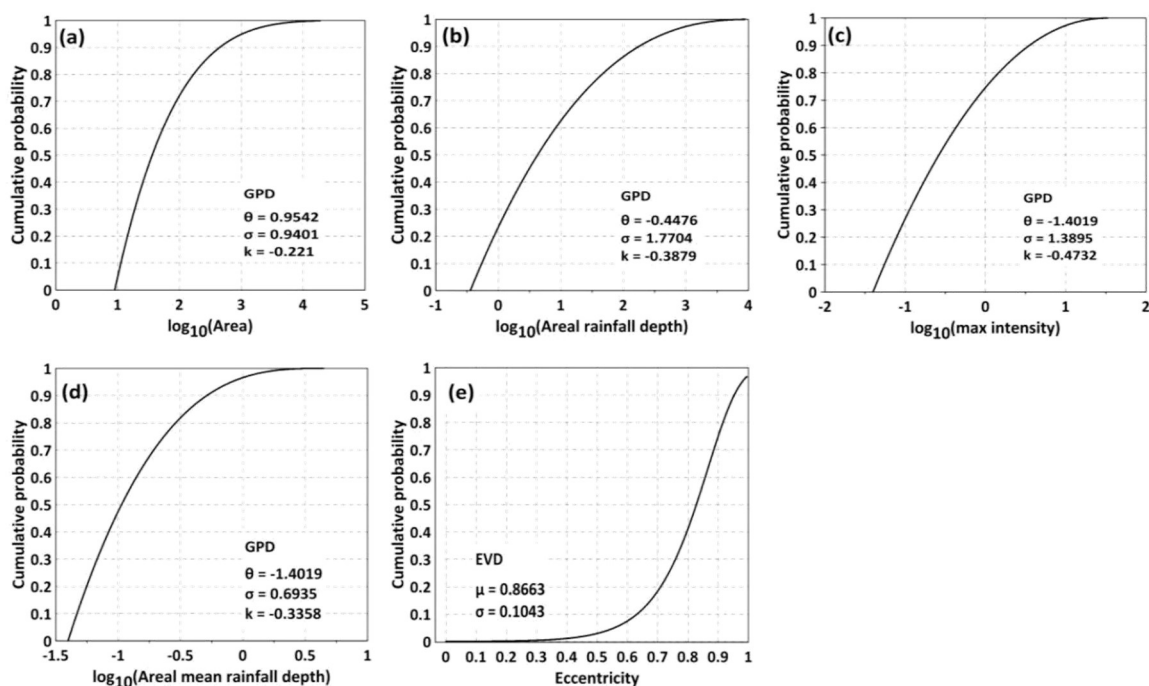
300 for $1 + k \frac{(x-\mu)}{\alpha}$, when $k > 0$, the generalized EVD is the Frchet distribution; $k < 0$ corresponds to the
 301 Weibull distribution; at the limit for $k = 0$, it is the Gumbel distribution.

302 A total of 1,107 rain cell trajectories were exploited. Histograms of their duration and motion
 303 vectors are presented in Figure 8. All rain cells held a mean duration of 40 minutes. For all the
 304 identified rain cells, the median value of their life cycles was 15 minutes, with an average moving
 305 speed of $11.59 \text{ m}\cdot\text{s}^{-1}$. The moving directions of the rain cells were consistently toward the direction of
 306 motion observed in the radar images.

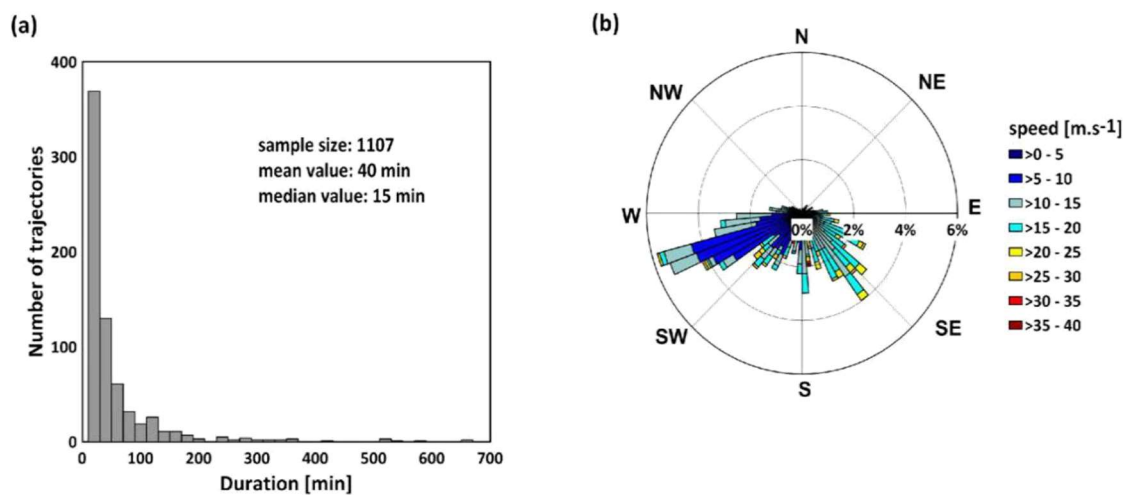


307

308 **Figure 6.** Histograms of \log_{10} -transformed rain cell properties: (a) area, (b) areal rainfall depth, (c)
 309 maximum intensity, (d) areal mean rainfall depth and property, (e) eccentricity for identified rain
 310 fields.



311 **Figure 7.** Cumulative curves of fitted probability density functions for \log_{10} -transformed rain cell
 312 properties: (a) area, (b) areal rainfall depth, (c) maximum intensity, (d) areal mean rainfall depth and
 313 property, (e) eccentricity.
 314



315 **Figure 8.** (a) Histograms of rain cell duration for identified rain fields, (b) wind rose plot of rain cell
 316 motion estimation result.
 317

318 These results were in agreement with the study of Barnolas et al. (2010), in which the structures
 319 of heavy rainfall events recorded in Catalonia, Spain were analyzed. However, the results differed
 320 from those of Karklinsky and Morin (2006), in which the area of identified rain cells in southern Israel
 321 was better fitted to the log-normal distribution. Statistical analysis of the rain cell properties suggests
 322 that the inner structures of the selected rainy days can be expressed by the EVD. This suggests that
 323 most rainfall events had a limited covering area with less intensity and short duration; rainfall events
 324 with a long duration had a large covering area and high intensity.

325 During the life cycle of a rainfall event, the physical and geometrical features of rain cells
 326 continually change. Three common stages reflect these variations: cumulus, mature, and dissipating
 327 (Byers and Braham Jr 1948). In fact, the stage changes of rain cells are not only associated with its
 328 internal growth and decay but also with outer rain cells (e.g., merging or splitting). In this study,
 329 seven rain cell life stages were confirmed by the RCIT algorithm; their definitions can be listed as
 330 follows (Figure 9):

- 331 a. Initial: A rain cell having no parent cell is termed an initial rain cell.
- 332 b. Tracking: A rain cell with only one parent cell and having no interaction with other rain
 333 cells during its life cycle is termed a tracking rain cell.
- 334 c. Merge: A rain cell with at least two parent cells is termed a merged rain cell.
- 335 d. Split: A rain cell with only one parent cell but at least two child cells is termed a split rain
 336 cell.
- 337 e. Dissipate: A rain cell with at least one parent cell but no child cells is termed a dissipate
 338 rain cell.
- 339 f. 5-minute life cycle: A rain cell with a life cycle of only 5 minutes.
- 340 g. Complex stage: A rain cell for which merging and splitting simultaneously exist during its
 341 life cycle is termed complex stage.

342 The number of rain cells with different stages over their life cycles is summarized in Table 4.
 343 Rain cells with "5 minutes life cycle" and "tracking" stages were dominant. The "merging",
 344 "splitting", and "complex stage" cells occurred in isolated cases, indicating a stable inner structure
 345 of the identified cells. For the cases of July 19, 2008 and July 26, 2008, the number of rain cells in the
 346 "tracking" stage was even greater, indicating that the rainfall events occurring on these two days had
 347 long durations. The number of rain cells in the "merging" and "splitting" stages was greater in the
 348 July 26, 2008 case. This suggests that there were more convective rainfall events on that day since rain
 349 cells merge or split more frequently under such conditions.

350 **Table 4.** Number of rain cells with different life stages, sorted by selected rain day.

Stages	May 26, 2007	July 19, 2008	July 26, 2008
Initial	158	350	471
Tracking	608	1270	1787
Merge	7	6	39
Split	1	2	5
Dissipate	152	346	434
5 minute life cycle	632	3148	929
Complex stage	0	0	1

351

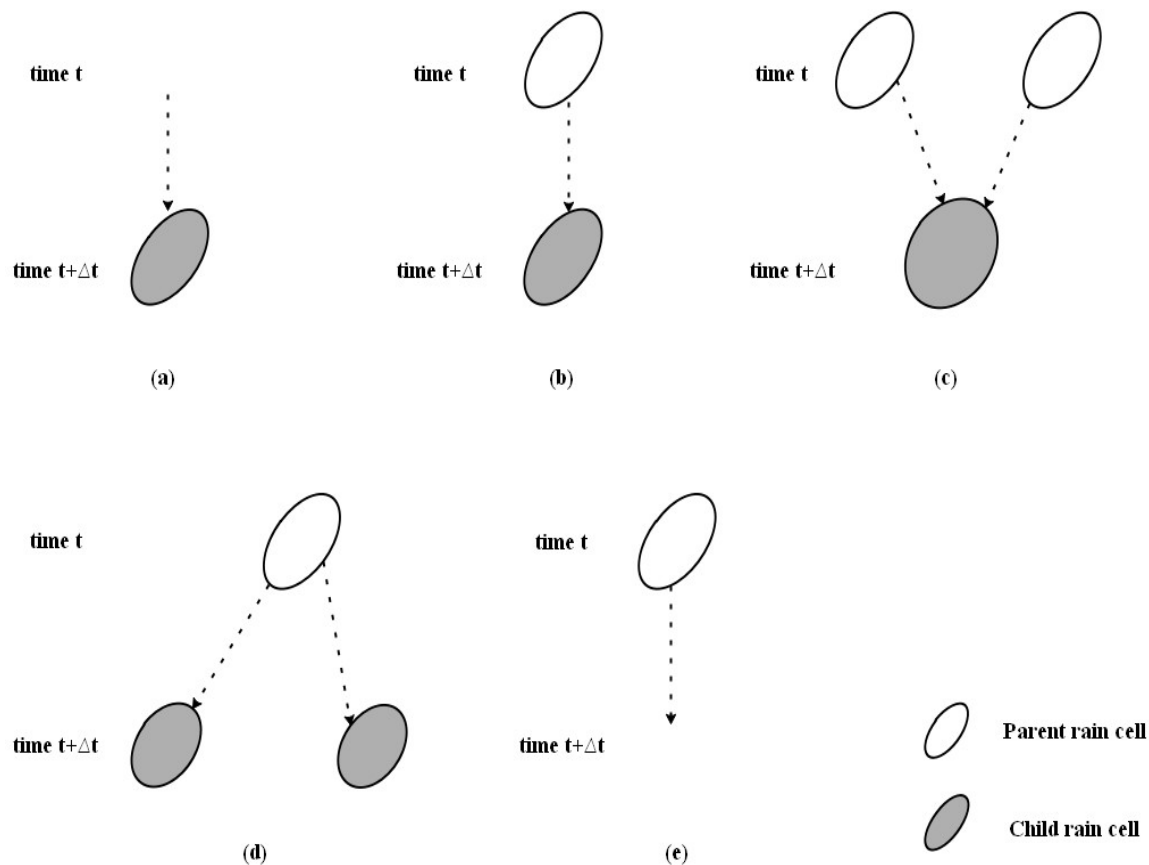


Figure 9. Stage definitions of rain cells: (a) initial, (b) tracking, (c) merging, (d) splitting, (e) dissipating.

352

353

354

355 5. Conclusion and Outlook

356 This study develops a new algorithm, RCIT, which utilizes the advantages of high resolution
357 weather radar data. The proposed algorithm provides the following improvements:

- 358 1. It uses the PIV method in rain cell motion estimation. Rain cell motion estimation by past
359 algorithms is mainly based on the maximum correlation coefficient method, which may
360 lead to nonconsecutive motion when the shape and volume of a rain cell change rapidly.
361 The PIV method avoids this situation.
- 362 2. A rain cell matching rule is proposed to discern the life cycle and stage change of rain cells.
363 Past algorithms focus mainly on the tracking of rain cells without merging and splitting,
364 when in fact rain cell stage variation is obvious over their life cycle, especially for
365 convective rainfall events. The proposed rain cell matching rule implemented in the RCIT
366 algorithm can easily and effectively discern the various stages of rain cells.

367 Two feature-based verification methods, SAL and geometric index, were used to test the
368 performance of the RCIT algorithm. It is shown that all verification indexes fall within in a reasonable
369 error range, confirming the good performance of the RCIT algorithm. Practical applications of the
370 RCIT algorithm in analyzing the inner structure of historical rainfall events that occurred in the NRW
371 are presented. This is the first time that the use of such a RCIT algorithm to depict the inner structures
372 of rainfall events in an urban region with a high population density has been presented. The results
373 show that the properties of rain cells in this region presented an EVD, indicating that the selected
374 rainfall events had a short duration with low intensity. Long duration events with high intensity are
375 rarely found and the stage changes of rain cells vary between events.

376 It should be noted that inputs for the proposed algorithm is not limited to radar data; other 2-D
377 remote sensing data will also be used as the algorithm inputs, suggesting the versatility of the

378 proposed algorithm. In future application, it is intended that this algorithm will analyze the spatial-
 379 temporal variation of rainfall in small regions; this will lead to the determination of rainfall inputs
 380 with proper spatial and temporal scales for hydro-meteorological applications. The proposed
 381 algorithm will also be applied to rainfall nowcasting, which will improve the foresight period of flash
 382 floods in mountainous and urban regions. In addition, the features of the rain cell output from this
 383 algorithm can be used in sensitivity analyses of urban runoff in relation to short-term rainfall events,
 384 which will improve flood forecasting precision in small-medium catchments.

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390 APPENDIX A

391 Structure–Amplitude–Location (SAL) and Geometric Index Verification Methods

392 Two feature-based verification methods, structure–amplitude–location (SAL) and geometric
 393 index, were applied to evaluate the performance of the RCIT algorithm and nowcasting methods. In
 394 SAL, term structure denotes the similarity in the shapes of modeled and observed rain cells; its value
 395 range is from -2 to 2. Amplitude denotes the similarity in the total precipitation values of modeled
 396 and observed rain cells; its value range is from -2 to 2. Location denotes the similarity of the center
 397 of mass for the modeled and observed rain cells; its value varies from 0 to 2. The accuracy of
 398 nowcasting methods can be evaluated based on the value of the three SAL components and a perfect
 399 nowcasting is confirmed by S, A, and L values of 0. More details on the SAL method can be found in
 400 Wernli et al. (2008).

401 Geometric index is a quantitative assessment method for the spatial patterns of rain cells
 402 (AghaKouchak et al. 2010). It compares the geometric features of modeled and observed rain cells via
 403 three indexes:

- 404 ● Connectivity index: This is defined to compare simulated rain cells with respect to a
 405 reference object (e.g., observed rain cells). Its value is calculated based on the number of
 406 rain cells (NC) and the total number of non-zero pixels or pixels above a given threshold
 407 (NP), as in Equation (8):

$$408 \quad C_{index} = 1 - \frac{NC-1}{\sqrt{NP+N}} \quad (8)$$

409 where C_{index} is the connectivity index, NP is the number of rainy pixels above a given
 410 threshold, and NC is the number of rain cells.

- 411 ● Shape index: This index is introduced to quantitatively describe the shape discrepancy of
 412 rain cells, as in Equation (9):

$$413 \quad S_{index} = 1 - \frac{P_{min}}{P} \quad (9)$$

414 where S_{index} is the single index, P_{min} is the theoretical minimum perimeter, and P is the
 415 actual perimeter of the rain cell.

- 416 ● Area index: This is defined to depict the dispersiveness between the modeled and
 417 observed rain cells. Its value is the ratio of the area of its convex hull (the boundary of the
 418 minimal convex set containing a finite set of points in the rain cell), as in Equation (10).

$$419 \quad A_{index} = 1 - \frac{A}{A_{Convex}} \quad (10)$$

420 where A is the area of the rain cell and A_{Convex} is the area of the convex hull.

421 APPENDIX B

422 Goodness of fit testing for fitted distributions of rain cell properties

423 The GOF test determines whether a data set is well fitted with a predefined distribution that
 424 gives the highest probability of producing the observed data. As such, a series of fit testing methods
 425 was developed, with the commonly applied tests as follows:

- 426 ● The K–S test is based on the empirical cumulative distribution function (ECDF). Given N
 427 ordered data points Y_1, Y_2, \dots, Y_n , their ECDF is defined as:

$$428 \quad E_N = \frac{n(i)}{N} \quad (11)$$

429 where $n(i)$ is the number of points less than Y_i and Y_j , which are ordered from the smallest
 430 to largest value. This is a step function that increases by $1/N$ at the value of each ordered
 431 data point. The K–S test was developed according to the following hypotheses: H_0 —the
 432 data follow a specified distribution; H_1 —the data do not follow the specified distribution.

- 433 ● AIC (Akaike, 1998) is based on the use of Kullback–Leible information as the discrepancy
 434 measure between the true distribution and the approximating distributions: $M_i =$
 435 $g_i(x, p_1, p_2, \dots, p_n)$. The AIC for the i th candidate distribution can be computed as:

$$436 \quad AIC = -2 \ln \prod(\theta) + 2p \quad (12)$$

437 where $\ln \prod(\theta)$ stands for the maximum log-likelihood of the sample of the dataset, p is the
 438 parameter's number of candidate distributions when the sample size n is small with
 439 respect to the number of the estimated parameter P_i . The smaller the value of AIC, the
 440 better fitting is the result for the candidate distribution.

- 441 ● BIC (Schwarz, 1978) serves as an asymptotic approximation to a transformation of the
 442 Bayesian posterior probability of a candidate model. It is based on the empirical log-
 443 likelihood and does not require the specification of priors. BIC is defined as

$$444 \quad BIC = -2 \ln \prod(\theta) + \ln(n)p \quad (13)$$

445 where the symbols are the same as those in Equation (12). A small value of BIC means
 446 that the candidate distribution fits well with the empirical distribution.

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528 precipitation forecasting using an object-based approach. *J Hydrol*, 483: 1–15

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542 box at a previous timepoint (red block), the MQD algorithm is applied to deduce the minimum
543 reflectivity differences, and grid points with minimum value in the next window box are
544 identified (blue blocks); step three: the solitary locations of reversed MQD value, Δx and Δy ,
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546 timepoint are calculated using the second polynomial function, and the global motion vectors
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548 Figure 3. Study area of North Rhine Westphalia (NRW) and its location in Germany. The main
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551 rain rate according to DWD standard Z–R relationship.

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555 rain days: (a) May 26, 2007; (b) July 19, 2008; (c) July 26, 2008.

556 Figure 6. Histograms of \log_{10} -transformed rain cell properties: (a) area, (b) areal rainfall depth, (c)
557 maximum intensity, (d) areal mean rainfall depth and property, (e) eccentricity for identified
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