Prediction of Holocene Mercury Accumulation Trends by Combining Palynological and Geochemical Records of Lake Sediments (Black Forest, Germany)

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Abstract

Forest vegetation plays a key role in the cycling of mercury (Hg) and organic matter (OM) in terrestrial ecosystems. Litterfall has been indicated as the major transport vector of atmospheric Hg to forest soils, which is eventually transported and stored in the sediments of forest lakes. Hence, it is important to understand how changes in forest vegetation affect Hg in soil and its biogeochemical cycling in lake systems. We investigated the pollen records and the geochemical compositions of sediments from two lakes (Schurmsee and Glaswaldsee) in the Black Forest (Germany) to evaluate whether long-term shifts in forest vegetation induced by climate or land use influenced Hg accumulation in the lakes. We were particularly interested to determine whether coniferous forests were associated with a larger export of Hg to aquatic systems than deciduous forests. Principal components analysis followed by principal component regression enabled us to describe the evolution of the weight of the latent processes determining the accumulation of Hg over time. Our results emphasize that the in-lake uptake of Hg during warm climate periods, soil erosion after deforestation and emissions from mining and other human activities triggered changes in Hg accumulation during the Holocene stronger than the changes caused by forest vegetation alone.
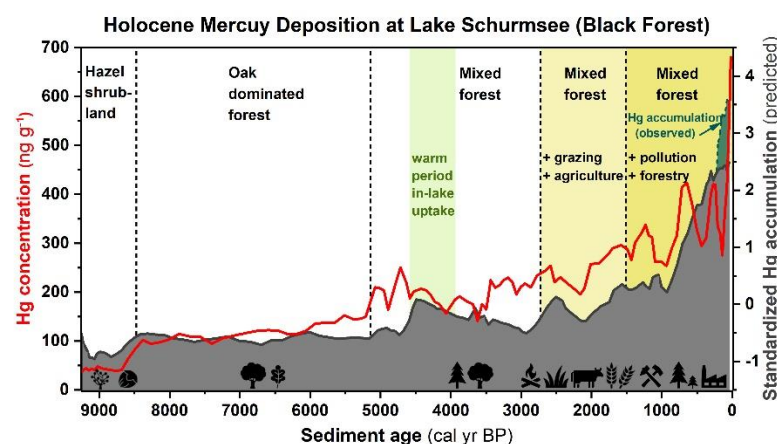
1. Introduction

Mercury is a pollutant of concern due to its toxicity and high bio-magnification potential in aquatic systems; because Hg is organophilic, OM is a key component to determine the retention and transport of Hg in the environment [1,2]. Widely observed increases in Hg concentrations and accumulation rates in lake sediments deposited during the past century are commonly accepted as records of increasing anthropogenic input of Hg to the atmosphere and atmospheric deposition of Hg to lake sediments [3–6]. Lake sediments accumulate material from the entire catchment and thus can be understood as a sink and an active pipe for OM and many major and trace elements, including Hg [7,8]. Because new material is deposited on top of older deposits, these sediments are also natural archives for temporal changes in trace element loadings [9–11]. Most studies of Hg in lake sediments have focused on correcting Hg accumulation in sediments for catchment effects to convert Hg accumulation rates into atmospheric deposition rates [5,12,13]. The effects of the evolution of landscapes on Hg transport and the accumulation of this element to and within aquatic systems have been considered to a much lesser extent. It is well established that the vegetation type influences the cycling of OM and trace element transport in several ways. First, the presence of vegetation stabilizes catchment soils and reduces the transport of particulate materials to surface waters [14]. Second, vegetation has an influence on the amount of litter that is produced, e.g., deciduous trees produce more litter than coniferous trees, which affects the trace element concentrations in the released dissolved organic matter (DOM) [15,16]. Third, vegetation and OM quality affect soil development, e.g., deciduous litter promotes the presence of invertebrates and soil mixing, which results in longer contact times between OM and soil mineral particles [17]. Fourth, differences in plant species compositions affect the interception area, which in turn affects the sequestration of atmospherically deposited trace elements [1,18]. Fifth, litter from different vegetation types decomposes at different rates, depending, e.g., on the litter chemical composition [19,20]. This affects the production of DOM and the amount of trace elements bound to each unit of OM so that more decomposed OM has higher Hg concentrations. Some studies have shown that Hg concentrations are higher in the litter of coniferous forests than deciduous forests [21,22]. In contrast, others studies have demonstrated that Hg concentrations in coniferous litter may be lower than those in deciduous litter [1,23,24], which might be due to varying conditions of atmospheric Hg concentrations or exposure. Jiskra et al. [25] indicated, with Hg isotope analyses in boreal soils, that 90% of Hg in forest soil is deposited by litter. While the vegetation acts as a short-term sink for atmospheric Hg, forest soils can be regarded as long-term sinks [26]. Drenner et al. [27] linked Hg levels in freshwater fish to the percentage of the landscape covered by coniferous forests. This highlights the significant effect of vegetation on the transfer of Hg to aquatic systems. Large natural

variations in the Hg content in lake sediments have been observed, suggesting natural site history, soil typ, morphology, hydrology, and climate variability [28–31].

One option to reduce the effect of site history, soil type, morphology and hydrology differences is to investigate the long-term effects of vegetation change and landscape development in a single system using long sediment records [32]. In this study, we analysed lake sediment records from two small headwater lakes, Schurmsee and Glaswaldsee in southwest Germany. The two lake sediment cores extend back to ~9500 yr BP and reflect two major vegetation changes in the Holocene. During the time period from 9500 to 2500 yr BP the atmospheric Hg deposition in central Europe was relatively constant [33,34]. Therefore, the two lake systems can be used as natural archives to reconstruct how changes between diverse vegetation cover types have affected the accumulation of Hg in the sediment from a long-term perspective. Considering the results of a preliminary work in this region by Rydberg et al. [32], we expected a pronounced variation in the Hg concentration in the lake sediments after palynologically evidenced changes in vegetation, and a significant increase in human impacts on the sediment geochemical composition from the Bronze Age onwards.

The main objective of this research was to combine the pollen-derived vegetation history and the geochemical composition of sediment records from the two lakes to assess how vegetation shifts affect the accumulation of Hg in lakes. Using multivariate statistical methods, we attempted to determine the main factors that controlled Hg deposition in lake sediments over a long-time scale (from Mesolithic to Modern Age) and to quantify the relative effects of these factors on Hg accumulation. This provides an opportunity to determine whether the observed changes in Hg accumulation were primarily caused by shifts in vegetation or whether other catchment-specific factors overwrote the influence of vegetation.

2. Materials and Methods

2.1 Site Description

The two studied lakes, Schurmsee (795 m a.s.l., 48°36′47′′N, 8°19′09′′E) and Glaswaldsee (839 m a.s.l., 48°25′33′′N, 8°15′41′′E), are small headwater lakes situated 23 km apart in the northern Black Forest, southwest Germany (Figure 1). The anthropogenic pressure on the lakes is currently relatively low. There are no settlements in the catchments, and only small forest tracks surround the study sites. The lakes have similar maximum water depths of 11-13 m, surface areas of 1.5 ha and 2.9 ha and lake-to-catchment-area ratios of approximately 1:34 and 1:18, respectively. The catchment topography was formed in a cirque excavated by glaciers, so the cupped sections of both lakes are relatively steep. The shoreline of Schurmsee is lined with a floating mat of *Sphagnum*. Surrounding

soils at both sites are dominated by acidified podzolic soils from Triassic sandstone rubble [35]. The climate is sub-oceanic, with decreasing temperatures and increasing precipitation as altitude increases. The annual average temperature, e.g., in Freudenstadt (732 m a.s.l., 48°27'4''N, 8°24'3''E), is 7°C, the mean annual precipitation is 1680 mm, and in that area, there are approximately 108 frost days (T_{min} = <0°C) per year [36]. Water budget measurements from 1985 to 1992 in the catchment at nearby Herrenwiesersee yielded a mean annual precipitation of 1931 mm [37]. The current lake water pH is 4.2-4.9 in the water column at both lakes. Depending on the water depth, the dissolved organic carbon concentration ranges from 6-15 mg L^{-1} at Schurmsee and from 15-19 mg L^{-1} at Glaswaldsee. The dissolved inorganic carbon is <0.5 mg L^{-1} and considered negligible. The present-day vegetation of the catchment area at both sites consists of managed coniferous forests dominated by *Picea abies* (Norway spruce). Pollen analysis of local sediment cores revealed that, similar to many other areas in Europe, several vegetation changes occurred at the study site during the Holocene [38,39].
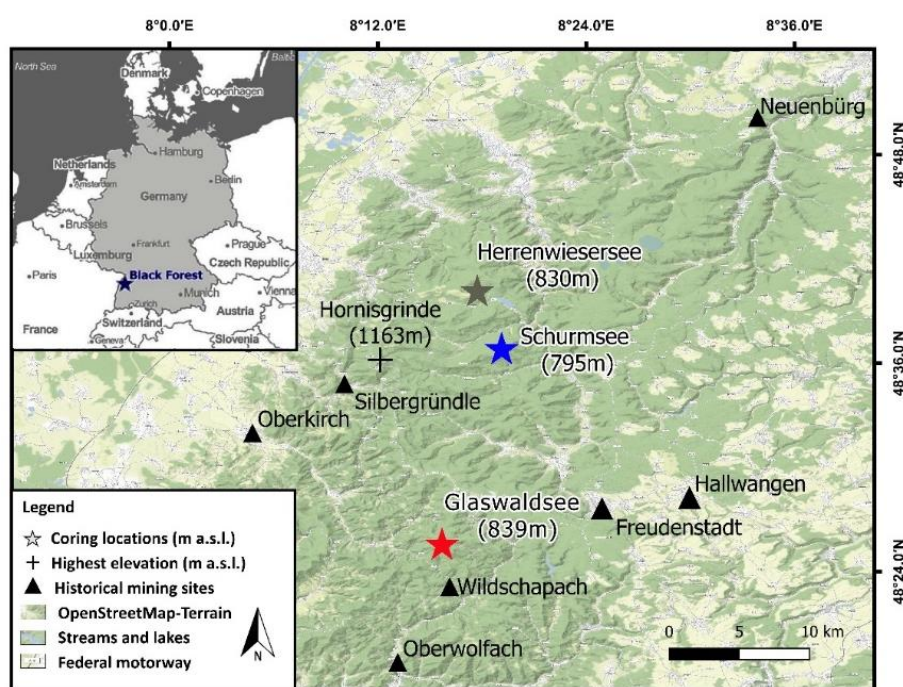
Figure 1: Map of the northern Black Forest showing the coring locations at Herrenwiesersee (grey star, Rydberg et al. [32]), Schurmsee (blue star), and Glaswaldsee (red star), as well as the topography, former mining sites and present-day land cover. Inset: localization of the study area in central Europe.

2.1.1 Coring

In 2006, two long cores were collected in the centre of the lakes at maximum water depth using a modified Livingstone-corer [39,40]. The long cores have a diameter of 5 cm and lengths of 526 cm at Schurmsee and 175 cm at Glaswaldsee. The cores were sub-sampled with a high resolution of 1 cm for pollen analysis and with a lower resolution of 4 cm and 2 cm for geochemical analysis, at Schurmsee and Glaswaldsee, respectively. For statistical reasons, the resolution of the pollen data used in this paper was reduced to the resolution of geochemical data for both cores. Data for the upper 28 cm were not available at Schurmsee. Therefore, a short sediment core was taken in 2017 using a UWITEC gravity corer. At Schurmsee, pollen data begin at a depth of 28 cm. According to the pollen spectra preserved by mosses from the shore of Herrenwiesersee [41], the pollen values for the statistical approaches were considered constant over the upper 28 cm at Schurmsee. A detailed description of the sample preparation and palynological analysis methods is given in Rösch and Lechterbeck [42]. The samples for the geochemical investigations were stored at -20°C in the dark until analysis. All core samples were oven dried at <40°C and ground using a Retsch agate ball mill.

2.1.2 Chronology

The sediment chronology was established by a ^{14}C-analysis, using 17 bulk sediment samples for Schurmsee and 16 bulk sediment samples for Glaswaldsee (Supplementary Materials, Table SM-1 and SM-2). Calibrated ages were obtained using the InterCal13.14C curve [43], and age-depth modelling was conducted with CLAM 2.2 using the smooth spline function [44]. One dating point of the Glaswaldsee core at 55 cm depth was excluded from the age-depth model because inclusion of the point resulted in an unrealistic sediment accumulation rate in this section of the profile. Samples below 400 cm at Schurmsee were not included in this study due to missing ^{14}C data and indicated multiple impacts of post-glacial erosion. Based on the age-depth model, the deepest sediment layers at both lakes were dated to ~9500 calibrated (CAL) yr BP. The average error for the age-depth estimates was 110 yr at Schurmsee and 183 yr at Glaswaldsee (Figure SM-1 and SM-2).

2.2 Chemical Analysis

2.2.1 Carbon, Nitrogen and Sulphur

Total carbon (C), nitrogen (N) and sulphur (S) were analysed using a EuroEA 3000 elemental analyser (Hekatech GmbH, Germany). The ratios of C:N are an indicator of organic matter sources [45]. The quality of the analysis was ensured by the inclusion of five certified reference materials (CRM). Two of the CRMs were pure chemical substances, i.e., sulphanilamide (C = 41.75 ± 0.17%; N = 16.26 ± 0.22%; S = 18.64 ± 0.18%) and BBOT (2.5-Bis(5-tert-butyl-benzoxazol-2-yl)thiophene; C = 72.52%; N = 6.51%; S = 7.44%), while three represented natural samples, i.e., NIST 1515 apple leaves (N = 2.25 ±

0.19%), Canmet LKSD-4 lake sediment (C = 17.7 ± 0.8%; S = 0.99 ± 0.09%) and MOC soil standard C = 3.19 ± 0.07%). Average results for the SRMs were sulphanilamide (n = 60) C = 41.86 ± 0.42%, N = 16.47 ± 2.54% and S = 18.87 ± 2.71%; BBOT (n = 14) C = 72.42 ± 0.52%, N = 6.49 ± 3.08% and S = 7.17 ± 4.69%; NIST 1515 (n = 9) N = 2.34 ± 2.56%; LKSD-4 (n = 4) C = 18.88 ± 0.62% and S = 0.98 ± 1.53% and MOC (n = 14) C = 3.24 ± 2.52%. Triplicate analyses (n = 14) gave relative standard deviations (RSD) of 1.57%, 2.26% and 2.71% for C, N, and S, respectively. According to pre-measurements using the "Karbonat-Bombe" method [46], the carbonate concentrations in the sediment were negligible. Hence, the measured carbon concentrations can be considered equivalent to the total organic carbon (TOC).

2.2.2 Major and Trace Elements

Total concentrations of several major and trace elements, i.e., aluminium (Al) titanium (Ti), iron (Fe), manganese (Mn), copper (Cu), zinc (Zn), zirconium (Zr), and lead (Pb), were analysed by means of energy-dispersive X-ray fluorescence (ED-XRF). Repeated analysis of two CRM lake sediments (Canmet LKSD-4, IAEA-SL-1) and sample triplicates gave RSD less than 10% for the CRMs, Ti, Zn, Zr, Fe and Pb, 7-19% for Al, 4-18% for Mn, and 3-23% for Cu. A detailed explanation about the calibration method, accuracy and precision is given in Cheburkin and Shotyk [47].

2.2.3 Mercury

The Hg content in the lake sediments was analysed by thermal decomposition followed by pre-concentration of Hg on a gold trap and cold vapour atomic absorption spectrophotometry (CV-AAS) Hg detection using a Milestone DMA-80 direct mercury [48]. Mercury concentrations are reported per dry weight. To ensure the analytical quality, two CRM lake sediments, i.e., Canmet LKSD-4 (190 ± 17 ng g^{-1}) and IAEA-SL-1 (130 ± 50 ng g^{-1}), were included alongside the analysed samples. The average results for the CRMs were as follows: LKSD-4 (n = 25), Hg = 191 ± 7 ng g^{-1} and IAEA-SL-1 (n = 6), Hg = 125 ± 4 ng g^{-1}. Randomly selected samples were analysed as triplicates (n = 18). The RSD between sample replicates was 1.2% (range 0.1-6.1%), which ensures that the samples were homogeneous.

2.3 Sedimentation Rates

Sedimentation rates are naturally not steady over a long-term period such as 9500 yr BP (Figure SM-1 and SM-2). The mean sedimentation rate at Schurmsee was more than two-fold higher than that at Glaswaldsee. Nevertheless, the two sediment cores are suitable for comparison because their pollen histories show similar trends over time, which is a mandatory requirement for contrasting the geochemical evolution of both lake sediments in this study. The dry mass sediment accumulation

rate (g cm^{-2} yr^{-1}) was calculated as the product of the mean bulk density of a given sediment layer (g cm^{-3}) and the interval represented by the given section (cm), divided by the time duration (yr), estimated by the age-depth model. The Hg accumulation rate (i.e., observed) at each sediment layer (µg m^{-2} yr^{-1}) was calculated as the product of the Hg concentration (ng g^{-1}) and dry mass sediment accumulation rate (g cm^{-2} yr^{-1}).

2.4 Statistical Treatment

Principal component analysis (PCA) was performed by combining pollen records and geochemical data to identify the main changes in the geochemical composition and relate these changes to underlying environmental processes. The advantage of this approach is that the sediment core is viewed as one unit for which co-variations between different sediment variables are used to describe main patterns by a few significant components (CPs). Each component combines variables with similar downcore trends that are likely controlled by the same underlying processes, e.g., elements that are affected by erosion events in the catchment. The PCA was performed separately for both cores using the following 18 variables: Al, Ti, Fe, Mn, Cu, Zn, Zr, Pb C, N, C:N ratio, and hazel, oak, beech, fir, spruce, ribwort, and rye pollen. Moreover, a charcoal record that was only available from Glaswaldsee was included in the PCA of the Glaswaldsee study site. To avoid Hg influencing the principal components, Hg concentration was not used as a variable in the PCA. Prior to the analysis, all values were standardized (Z-scores) to avoid scaling effects and to obtain average centred distributions [49]. A varimax rotation was applied to maximize the loadings of the variables. Components with eigenvalues <1 were not considered and were sorted out.

In this study, we aimed to predict Hg accumulation based on environmental variables, including human activities, and to identify the primary effects influencing the Hg accumulation rates, without the influence of the bulk density, which may result in inaccurate observed Hg accumulation rate values. Based on the results of the PCA, a stepwise regression analysis (PCR) was performed using the Hg accumulation rate as the dependent variable and the scores of the extracted components as the predictors. The main advantage of using component scores rather than the original geochemical data in the regression analysis is that the rotated components represent new synthetic orthogonal variables, and the effects of these components within the regression analysis are additive. The record of the regression weights of the components with depth/time was used to evaluate the chronological variations of the controls on the Hg record [50]. All statistical calculations were performed within the statistic environment of R - version 3.3.3 [51].

3. Results

3.1 Vegetation Changes and Land Use Practices

The pollen records from Schurmsee and Glaswaldsee indicate that, after the post-glacial forest establishment, *Corylus avellana* (hazel) became the most dominant woody species until ~8500 yr BP (Figure 2); thereafter, hazel was replaced by *Quercus robur* (oak). A second major change occurred between 5700-5300 yr BP when *Abies alba* (fir) and *Fagus sylvatica* (beech) took over as the most dominant tree species. Considering the altitude of the lakes, this mixed forest was likely dominated by *Abies alba*, but *Fagus sylvatica* was also present in the catchment. Among the herbs and grasses, there were indicators for several types of cultivated landscapes. Clear traces of human impacts due to grazing were observed between 3100 yr BP and 2500 yr BP by *Plantago lanceolata* (ribwort). The pollen analysis of the local sediment cores showed that, similar to many other areas in Europe, several vegetation dynamics occurred during the Holocene. Human activity (i.e., land use near both lakes) was clearly indicated by *Secale cereale* (rye) after ~1200 yr BP. The pollen findings of *Picea abies* (spruce) were less than 5% until the Late Middle Ages. Due to the emergence of local forestry, spruce increased to ~20% during the 19th century [52] and currently represents the dominant tree species in the northern Black Forest. Based on the vegetation dynamics and land use practices, the sediment records were divided into five time periods (Figure 2, section I. to V.), e.g., for Schurmsee, these periods were prior to 8400, 8400-5100, 5100-2600, 2600-1600 and after 1600 yr BP. To improve the chronological comparability between the locations, a linear age-depth distribution was applied.
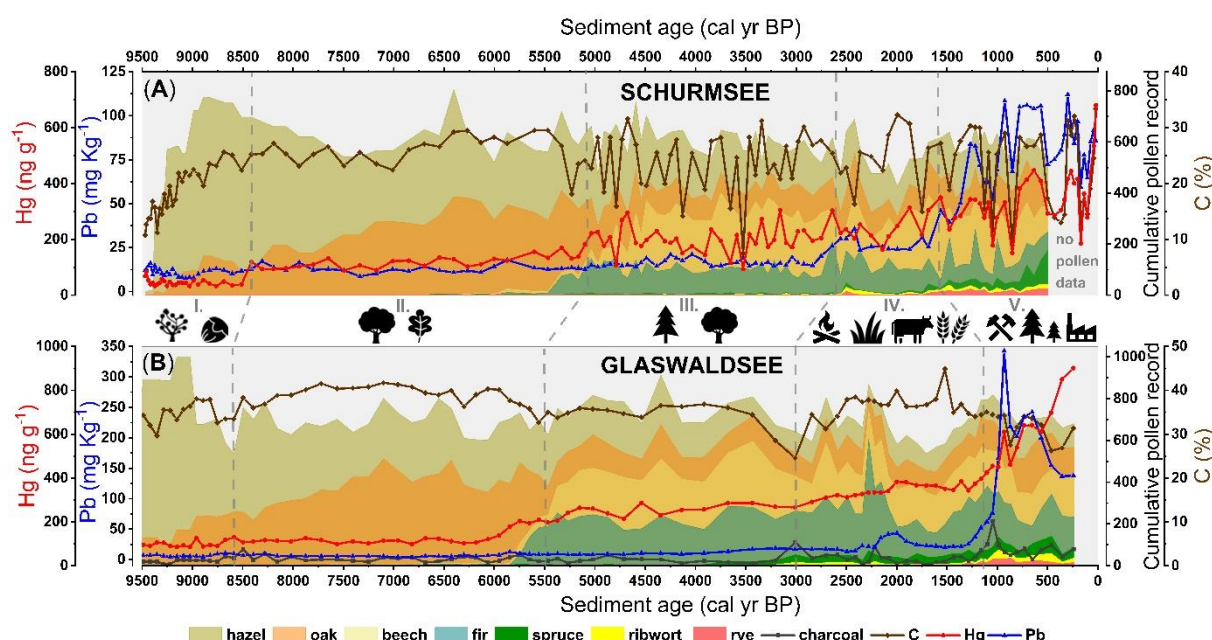
Figure 2: Cumulative historical pollen records for hazel, oak, beech, fir, spruce, ribwort and rye (coloured stacked area) and concentrations of C (brown line), Hg (red line), and Pb (blue line) for Schurmsee (A) and Glaswaldsee (B) by linear age-depth distribution. In addition, a charcoal record (grey line) is shown for Glaswaldsee, and indicators (black symbols) of the vegetation cover types and human-caused impacts are shown for both catchments.
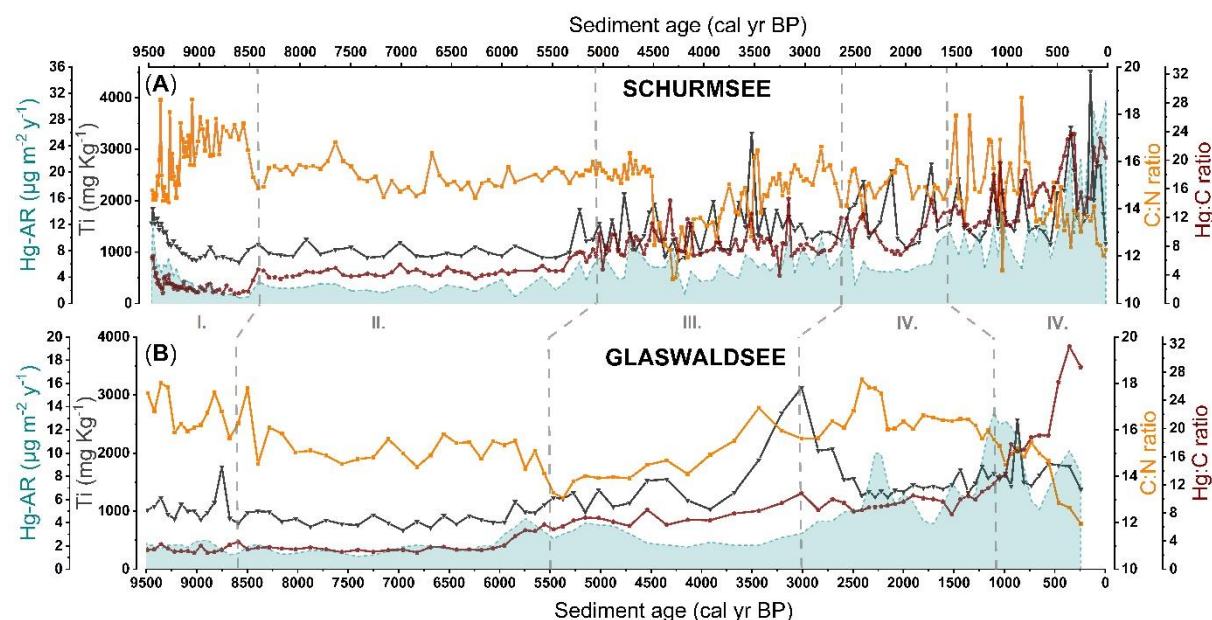
Figure 3: Titanium concentrations (dark-grey line) as indicators of mineral matter erosion, carbon/nitrogen ratios (orange line) as indicators of organic matter sources and observed mercury accumulation rates (grey-blue area), mercury/carbon ratios (dark red line) as an indicator of the interaction of atmospheric Hg fluxes and organic matter sources for Schurmsee (A) and Glaswaldsee (B) by linear age-depth distribution.

3.2 Holocene Records of Mercury and Lead Concentrations

Mercury concentrations varied from the oldest (section I.) to the earliest core layers (section V.) from 72 to 680 ng g^{-1} at Schurmsee and from 95 to 901 ng g^{-1} at Glaswaldsee. Schurmsee showed a sharp increase in Hg concentrations from 43 to 116 ng g^{-1} (Figure 2.A) when oak replaced hazel as the dominat tree species at ~8500 yr BP, while the Hg concentrations at Glaswaldsee increased from 117 to 130 ng g^{-1} during this first forest change. After the formation of a mixed forest with fir and beech at ~5100 yr BP at Schurmsee and ~5500 yr BP at Glaswaldsee, the Hg concentration increased to 182 ng g^{-1} and 209 ng g^{-1}, respectively. Since a mixed forest was established in the catchment of Schurmsee, the Hg record was more unstable, i.e., it was influenced by a number of erosion events, as indicated by rapid declines in the carbon concentrations (Figure 2), and simultaneous increases in the titanium concentrations (Figure 3). Between 3000 and 2600 yr BP, the Hg concentrations in both sediment cores started to increase simultaneously. During the same period, the first signals of grazing activities shown by ribworth pollen were observed. The Hg concentrations in these sediment layers (section IV.) reached up to 350 ng g^{-1} at Schumrsee and 344 ng g^{-1} at Glaswaldsee. However, the highest Hg concentrations in both cores were observed in the surface layers (section V.), which was temporally related to the Industrial Age. Hg concentrations reached up to 680 ng g^{-1} and 901 ng g^{-1}, respectively.

In comparison to the Hg records, the Pb records were relatively constant over core sections I. to III. (Figure 2), which showed a mean Pb concentration of 13.7 µg g^{-1} at Schurmsee and 8.3 µg g^{-1} at Glaswaldsee. The first significant increases in the Pb concentrations appeared in section IV. and were concurrent with other indicators (e.g., ribwort) of increasing human settlement activity. The strongest Pb signals during the High Medieval period peaked at ~930 yr BP in section V. and showed concentrations of 108 µg g^{-1} and 342 µg g^{-1}, respectively. Corresponding to the high phase of local mining during the Medieval period, the Pb concentrations decline significantly in the upper sediment layers at ~500 yr BP of both lakes.

3.3 Variation in Sediment Geochemistry Revealed from Principal Component Analysis

The PCA resulted in six significant (i.e., eigenvalues >1) components explaining 87% of the total variance in the palynological and geochemical data of both cores (Table 1). The first component (CP1), explaining 20% of the total variance at Schurmsee and 22% of the total variance at Glaswaldsee, was characterized by high loadings of C and N, high contrary loadings for the soil erosion indicators Zr, Al and Ti, and moderate positive loadings for Mn at Glaswaldsee. The observed human impact was explained by CP2 (25% and 19% of the total variance) and was indicated by high positive loadings of rye, ribwort, Pb, and spruce. Spruce pollen showed a loading of 0.90 at

Schurmsee, which was higher than the loading of 0.67 at Glaswaldsee. The charcoal record at Glaswaldsee resulted in a positive loading of 0.67 and indicated the anthropogenic influence within this component. The third component (CP3), which explained 18% of the total variance for both cores, was characterized by positive loadings for beech and fir, whereas hazel showed negative loadings. The fourth component (CP4), explaining 9% and 6% of the total variance at Schurmsee and Glaswaldsee, respectively, was characterized by high loadings for oak and negative loadings for Fe at Schurmsee. The fifth component (CP5), which explained 8% of the total variance at both cores, showed high loadings for the C:N ratio and a moderate inverse loading for Zn at Schurmsee. The sixth component (CP6) explained 7% and 14% of the total variance by high positive loadings of Cu at Schurmsee and high positive loadings of Cu, Zn, and Fe at Glaswaldsee.

Table 1: Factor loadings of six component (CPs) extracted by PCA using the pollen and geochemical element composition of the two Black Forest sediment cores: eigv. - eigenvalues, var. - percentage of proportional variance explained by each component. For reasons of comparability, no strict sorting according to loadings and eigenvalues was applied.

	SCHURMSEE							GLASWALDSEE					
proxies	CP1	CP2	CP3	CP4	CP5	CP6	proxies	CP1	CP2	CP3	CP4	CP5	CP6
C	**0,91**	0,15	0,11	0,17	-0,09	0,18	C	**-0,83**	-0,26	0,00	0,32	0,04	-0,23
N	**0,87**	0,25	0,20	0,11	0,26	0,17	N	**-0,74**	-0,14	0,07	0,28	-0,55	-0,03
Zr	**-0,83**	0,41	-0,05	-0,18	0,08	-0,09	Zr	**0,86**	0,19	0,02	0,06	0,10	0,25
Al	**-0,89**	0,27	0,24	0,01	0,08	-0,06	Al	**0,83**	0,08	0,31	-0,10	-0,08	-0,32
Ti	**-0,79**	0,42	0,33	-0,02	0,09	0,01	Ti	**0,80**	0,30	0,46	0,03	0,01	-0,13
							Mn	**0,59**	0,01	0,55	0,19	0,27	0,07
rye	-0,13	**0,95**	0,10	-0,09	0,04	-0,02	rye	0,12	**0,92**	0,08	-0,07	-0,10	0,01
ribwort	-0,14	**0,89**	0,25	-0,07	0,11	-0,04	ribwort	0,21	**0,86**	0,35	0,00	0,02	0,02
Pb	0,00	**0,91**	0,18	-0,10	0,05	-0,02	Pb	0,13	**0,88**	0,05	-0,20	-0,20	0,06
spruce	-0,13	**0,90**	-0,03	-0,17	0,27	0,02	spruce	0,15	**0,67**	0,59	-0,05	0,14	0,00
Mn	0,05	**-0,82**	-0,35	-0,19	0,12	0,12	charcoal	0,17	**0,67**	0,14	0,31	0,10	0,17
beech	0,03	0,12	**0,91**	-0,08	0,07	-0,06	beech	0,22	0,17	**0,85**	-0,24	-0,11	-0,18
fir	0,02	0,22	**0,83**	0,03	0,03	0,06	fir	0,17	0,20	**0,84**	-0,21	-0,01	-0,22
hazel	0,16	-0,36	**-0,65**	0,13	-0,21	-0,30	hazel	0,03	-0,29	**-0,80**	-0,23	0,25	0,16
oak	0,25	-0,03	-0,21	**0,83**	0,04	0,21	oak	-0,42	-0,08	-0,45	**0,68**	-0,27	0,05
Fe	-0,02	0,39	-0,39	**-0,60**	0,25	0,33							
C:N	0,13	-0,27	-0,33	0,11	**-0,84**	0,00	C:N	0,08	-0,16	-0,14	-0,05	**0,91**	-0,25
Zn	0,21	-0,12	-0,35	0,43	**0,55**	0,40							
Cu	0,37	-0,16	0,17	0,13	0,05	**0,80**	Cu	-0,13	0,04	-0,16	0,00	-0,22	**0,90**
							Zn	0,04	0,01	-0,24	0,16	-0,19	**0,86**
							Fe	0,38	0,20	-0,03	-0,17	0,23	**0,79**
eigv.	**3,15**	**4,04**	**2,87**	**1,50**	**1,29**	**1,12**	eigv.	**4,19**	**3,70**	**3,48**	**1,05**	**1,59**	**2,62**
var.	**20**	**25**	**18**	**9**	**8**	**7**	var.	**22**	**19**	**18**	**6**	**8**	**14**

3.4 Prediction of Factors Controlling Mercury Accumulation Using Stepwise Regression Analysis

Here, the aim was to predict Hg accumulation rates based on sediment variables, including indicators of human activities, to identify the primary effects causing Hg accumulation. In addition, we were interested in whether coniferous forests led to a larger export of Hg to aquatic systems than

deciduous forests. Table 2 presents the results of four (i.e., model 1a, 1b, 2a, and 2b) stepwise regression analyses (PCRs), which were performed using the Hg accumulation rate as the dependent variable and the scores of the extracted components as predictor variables. Regression model 1a implies that the score values from all six components were significant enough and were therefore used to explain the Hg accumulation at Schurmsee. Model 2a for the Glaswaldsee core shows that only three of the six components needed to be retained for the PCR to predict Hg accumulation. The regression models using the PCA variables suggest that Hg accumulation was predicted with a high accuracy at both Schurmsee (0.88, error ± 0.49) and Glaswaldsee (0.82, error ± 0.59). To assess the relevance of coniferous tree species on Hg accumulation, we excluded fir and spruce in a second step of the models. After excluding the conifers, shown here by the regression models 1b and 2b, the Hg accumulation predictions showed slightly lower, but still high correlation coefficients for Schurmsee (0.87, error ± 0.50) and Glaswaldsee (0.80, error ± 0.62). Because the analyses were performed with standardized variables, the regression coefficients were a measure of the weight of each factor on the variation of Hg accumulation over the last 9500 years in both lakes. In the models 1a and 2a, the factors weight decreases in the order of CP2 > CP5 > CP4 > CP1 > CP3 > CP6 at Schurmsee and CP2 > CP3 > CP5 at Glaswaldsee. The components CP2, CP3, and CP5 showed a positive relationship with Hg accumulation, while CP1 and CP4 implied a negative effect on Hg accumulation. In addition, CP6 signified a moderate positive effect on Hg accumulation at Schurmsee.

Table 2: Stepwise regression analysis of CPs on Hg accumulation in the two sediment cores: components involved in the variations of Hg accumulation, regression coefficients of the significant CPs explaining Hg accumulation, t- and p-values, correlation coefficients (R) and errors for the four regression models; models 1a and 2a included all variables; models 1b and 2b excluded pollen data of coniferous tree species.

	CP1	CP2	CP3	CP4	CP5	CP6	R	error
model 1a SCHURMSEE								
reg. coef.	-0,20	0,73	0,11	-0,26	0,33	0,08	**0,88**	0,49
t-value	-4,55	**16,83**	**2,47**	-6,02	**7,50**	**1,90**		
p-value	<0,001	<0,001	<0,05	<0,001	<0,001	<0,1		
model 1b SCHURMSEE								
reg. coef.	-0,21	0,67	0,36	-0,07	0,37	-	**0,87**	0,50
t-value	-4,59	**14,90**	**8,16**	-1,62	**8,25**	-		
p-value	<0,001	<0,001	<0,001	<0,1	<0,001	-		
model 2a GLASWALDSEE								
reg. coef.	-	0,68	0,44	-	0,11	-	**0,82**	0,59
t-value	-	**10,81**	**6,92**	-	**1,74**	-		
p-value	-	<0,001	<0,001	-	<0,1	-		
model 2b GLASWALDSEE								
reg. coef.	-	0,70	0,09	-0,17	-	-	**0,80**	0,62
t-value	-	**10,50**	**5,06**	-2,55	-	-		
p-value	-	<0,001	<0,001	<0,05	-	-		

4. Discussion

4.1 Evolution of Mercury Concentration and Accumulation

The sedimentary history of both lakes reflects strong environmental changes in the lake catchments over the past 9500 years. The Hg concentrations showed similar trends during the Holocene and increased from the oldest layers to the surface layers at Schurmsee (72-680 ng g^{-1}) and Glaswaldsee (95-901 ng g^{-1}) by factors of 9.4 and 9.5, respectively. Large ranges of Hg concentrations within lake sediment cores have been reported, e.g., from lakes in Sweden (~100-900 ng g^{-1}) [53], the United Kingdom (~20-1700 ng g^{-1}) [54] and Switzerland (~20-700 ng g^{-1}) [34]. Moreover, the ranges of Hg concentrations observed in the two cores are similar to those found in Herrenwiesersee [32]. In the earliest layer at Herrenwiesersee (1566 yr BP), the reported Hg concentration was 316 ng g^{-1}. This Hg concentration is slightly lower than the Hg concentrations found in comparable sediment layers in our study, showing 350 ng g^{-1} (1564 yr BP) at Schurmsee and 366 ng g^{-1} (1606 yr BP) at Glaswaldsee.

Long-term records encompassing Hg deposition over several thousands of years suggest a natural variability in the background values comparable to the recent anthropogenic forcing of the atmospheric Hg cycle [28,55,56]. Potential factors controlling pre-industrial Hg variations in lake sediments are natural changes in atmospheric fluxes. Long-term changes of Hg concentrations in polar ice, e.g., have been attributed to varying marine emissions coupled to oceanic productivity changes [57], whereas Guevara et al. [12] found a direct relationship between changing lake sediment Hg concentrations and volcanic eruptions in their study area. However, clear signals of Holocene volcanic eruptions could not be indicated in the two studied lake sediments. Besides the atmosphere, lake catchments are main Hg sources due to erosion of mineral particles or OM [58]. The amount of Hg transported from a catchment into adjacent aquatic systems is regulated by fluxes of terrestrial OM [59–62]. Hence, the Hg accumulation rates in lake sediments will be highly dependent on the export of Hg from catchment soils. In addition, the concentration will also be affected by, e.g., the influx of mineral matter relative to OM [63], the degradation state of OM [21], and the infilling of sediment with time that might affect sediment focusing [64].

In the two studied lakes, the terrestrial part of the catchment is considerably larger than the lake areas, but the ratio of lake surface area to catchment area is much higher at Schurmsee than at Glaswaldsee, which may have an import effect on the amount of OM transported to the lake sediments and their drivers (i.e., climate, vegetation). At Schurmsee, the recent Hg accumulation rate in the upper lake sediment layer was 31 µg m^{-2} yr^{-1} (Figure 3), which is significantly higher than the present-day measurements of regional wet deposition, e.g., 21 µg m^{-2} yr^{-1} at the northern Black

Forest [65], indicating that a significant portion of the Hg input to the lake sediment was derived from the catchment runoff. However, the observed Hg accumulation in the sediment at Glaswaldsee peaked at 15 µg m^{-2} yr^{-1} during the High Medieval mining period and showed a value of 8 µg m^{-2} yr^{-1} in the upper lake sediment layer.

According to Hermanns and Biester [56], the precipitation-induced mobilization of Hg-OM complexes from catchment areas is the main cause of the transport of Hg from terrestrial to aquatic ecosystems, as indicated by the similar Hg and C concentrations trends at Schurmsee, especially in section III. to V. (Figure 2.A). In addition, the catchment morphometry, e.g., differences in slope gradients and, lake forms, influences the concentration of Hg in the lake sediments [66]. Furthermore, increases in the Hg export from catchment soils can be caused by variations in the input of Hg by atmospheric deposition, e.g., due to increases in industrial activities, and the connectivity between terrestrial and aquatic pools, and decreases in the long-term retention or reemission of Hg [32]. All these factors can also potentially affect the Hg accumulation in sediments, even though most of these parameters are similar at both studied sites due to their geographic nearness and similar altitudes. Therefore, a multivariate statistical approach was considered.

4.2 Environmental Processes Controlling Sediment Geochemistry

The diverse palynological and chemical compositions of the sediments of Schurmsee and Glaswaldsee are represented by six components (Table 1). The CP1 of the PCA is characterized by large loadings of lithogenic elements (Zr, Al, Mn) and contrasting loadings of organic matter (C, N). The source-rock at both lakes is mainly Triassic sandstone, which is usually rich in quartz and partly rich in clay minerals, indicating that the high Al values in the sediments are geogenic. CP1 may represent a signal of local mineral fluxes due to soil erosion from the relatively steep slopes surrounding the lakes. Increased mineral matter fluxes (e.g., Ti) lead to the dilution of the OM in the sediments, reflected by the contrasting loadings of total carbon. CP2 represents the following signals of human activities in chronological order, starting in the northern Black Forest in the Early Iron Age: grazing > agriculture > mining [39,52,67]. With the emergence of the first rural cultures with livestock farming, anthropogenic disturbances and changes in the natural vegetation cover became increasingly common. While the Iron Age deforestation at ~2600 yr BP in the region was already high, the Medieval deforestation maximum at ~800 yr BP was two-fold that of the Iron Age and much more extended than modern deforestation [68]. CP3 represents the decline of an oak-dominated forest followed by the evolution of a mixed forest dominated by fir and beech. At Schurmsee, fir was distributed ca. 500 years earlier than beech. This forest change was most likely driven by a change in climatic conditions. According to Rösch [39], there are only weak traces of Neolithic human impacts in the northern Black Forest. Furthermore, there is no evidence for human settlements and land use

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for at least 2000 years after the occurrence of beech and fir at ~5000 yr BP. CP4 represents the variation of an oak-dominated forest and shows a contrasting Fe loading for Schurmsee. CP5 and CP6 shared variables related to the formation processes of autochthonous versus allochthonous OM and affiliated trace elements such as Cu and Zn. Moreover, Cu is an important co-factor in plastocyanin, a protein essential for the algae photosynthesis and is taken up by algae mainly as Cu^{2+} [69]. CP5 may have particularly a high importance because changes in aquatic productivity can be responsible for variations of Hg deposition in the sediments due to increased Hg scavenging by organic particles and the export of Hg to the sediment [70–72]. On the other hand, allochthonous organic matter from organic-rich eroded soil from catchments is generally older and more degraded than autochthonous OM from algae and may lead to higher Hg fluxes from the catchment to the lake because of the higher Hg:C-ratio [25,32,73].

4.3 Chronology of Factors Controlling Mercury Accumulation

While the environmental processes controlling the sediment geochemistry were previously represented by six components for the whole core length, the intensity (i.e., relative importance) of the different processes controlling Hg accumulation is not steady over time. The cumulative effects of the factors for each lake sediment layer are shown in Figure 4. Based on the classification by vegetation cover type and human impacts, the two sediment records can again be divided into five time periods (section I. to V.). In Figure 4, only PCR models 1a and 2a are considered. Regarding the type or environmental meaning of the variables forming the CPs within the PCR, the CPs are labelled follows: CP1 refers to soil erosion and contains opposing elements of minerals and OM. CP2 combines human impact indicators, e.g., rye for agriculture and Pb for mining activities. CP3 is associated with the evolution of mixed mountain forests (i.e., the presence of beech and fir). CP4 is linked to the presence of oak trees. CP5 describes the effect of the C:N ratio, quantifying the influence on Hg accumulation either from autochthonous or allochthonous OM.

Due to an additional contribution of zinc (Zn) in CP5 in the PCA at Schurmsee, which resulted in negative loadings for the C:N ratio, the bar representation for CP5 at Schurmsee is opposite of that for CP5 at Glaswaldsee. CP6 refers to the concentration of Cu, which is only relevant for model 1a.
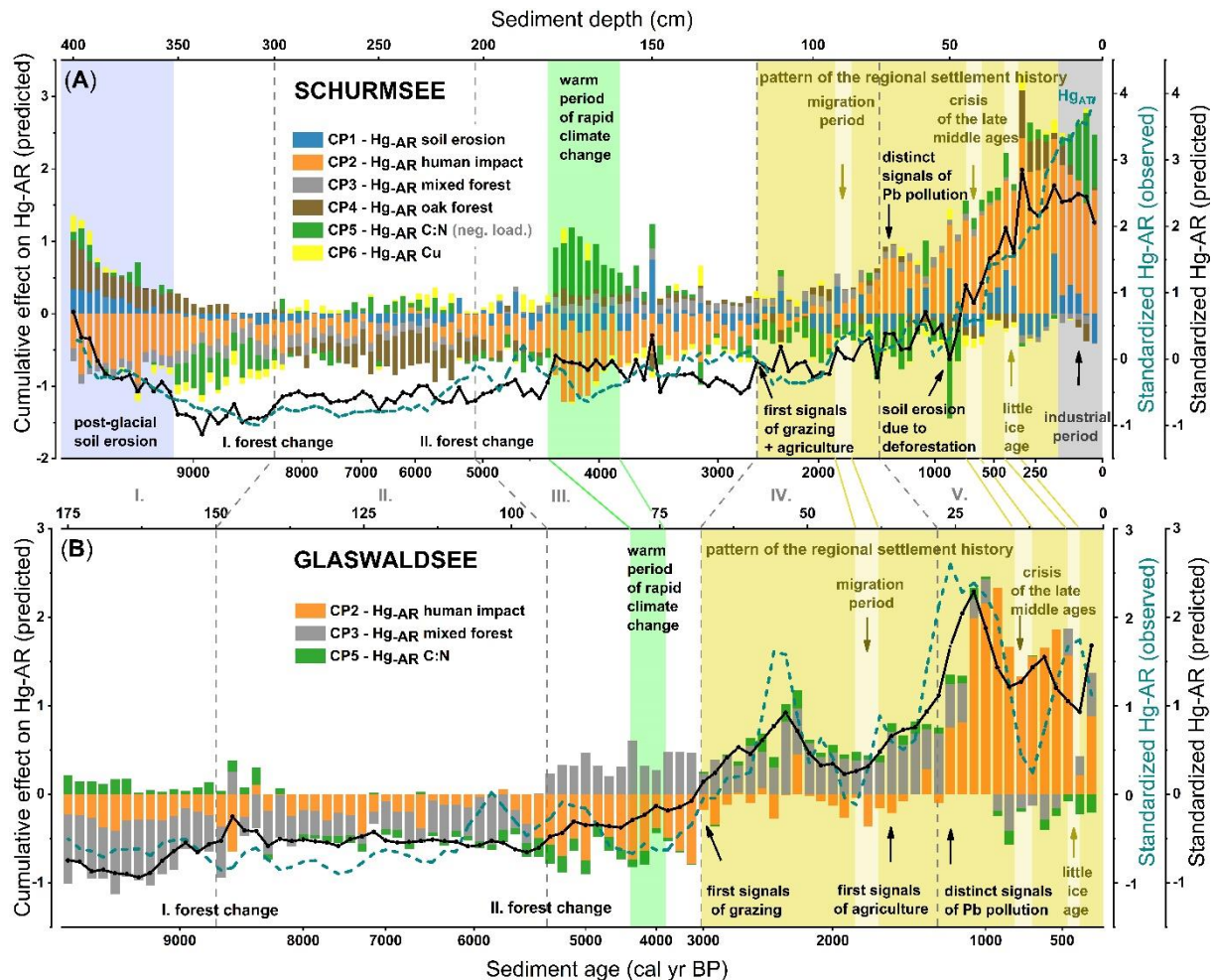
Figure 4: Chronology of the cumulative effects (coloured stacked bars) of the significant factors (CP-Hg-AR) involved in the variations of Hg accumulation (Hg-AR) at Schurmsee (A) and Glaswaldsee (B) by a lake-specific age-depth distribution. The patterns of climate and local settlement history, records of observed standardized Hg-AR (grey-blue dashed line) and predicted standardized Hg-AR (black line) were obtained by PCR models 1a and 2a.

Declining Soil Erosion in an Establishing Hazel-Dominated Shrubland (Section I.)

Intensive soil erosion was observed at Schurmsee until ~9200 yr BP and led to higher uncertainties in the age model and unrealistic Hg accumulation rates (Figure SM-1). After this period of post-glacial soil erosion, the low Hg accumulation is related to moderate inputs of mineral matter due to a higher erosion of OM under a vegetation cover of mainly hazel. The absences of human influences and mixed forest stands are clearly shown by the negative effects of the CP2 and CP3 factors, respectively. The concentrations of C are lower under the hazel-dominated forest than under the later oak-dominated forest, which is related to a lower input of OM from hazel stands. Furthermore, the mean C:N ratios between 16.7 and 16.9 indicate the OM was mainly from allochthonous sources in the earliest core section of both lakes [56]. Consistently, the observed and predicted Hg accumulations show similar courses, even during the post-glacial erosion period.

Steady Fluxes of Soil Organic Matter in an Oak-Dominated Forest (Section II.)

The predicted Hg accumulation rates increase moderately after the first forest change from hazel to oak. CP4 reflects the occurrence of an oak-dominated forest in this section. Both sediment cores show a rapid increase in Hg concentrations (Figure 2) and predicted Hg accumulation rates at the beginning of this period, although if CP4 is considered alone, it has a negative effect on Hg accumulation (Table 2). This result is consistent with the accumulation of a thick organic soil horizon in oak forest stands, which reduces soil erosion, i.e., decreases dilution by mineral matter in the sediment, and increases the flux of DOM-bound Hg to lake sediments [32]. Moreover, compared to the shrubby vegetation under hazel, oak trees develop a denser canopy, which could result in an increased interception of atmospheric Hg and reduce the evasion of volatile Hg [74–76]. At Schurmsee, the effect of CP4 and CP6 on Hg accumulation changed from negative to positive, which is likely related to a higher production of autochthonous OM by algae. This transformation can also be seen for CP5 at Glaswaldsee and during 7800-5700 yr BP at Herrenwiesersee [32]. According to a review of natural archives from central Europe by Kalis et al. [77], the Early Atlantic (9000 to 7500 yr BP) was a period characterized by negligible human impacts and relatively stable environmental conditions within the Holocene climatic optimum. Mayer and Schwark [78] have shown, with a stable isotope record obtained from the lacustrine sediment core at Steißlingersee (southwest Germany), an abrupt increase in $\delta^{13}C_{org}$ during the Early Atlantic, which was explained as a negative water balance signal due to the increased evaporation and temperatures during dry climatic conditions and probably caused the first forest change in the catchments of the studied areas. Eventually, reduced precipitation, decreases in mineral matter fluxes from the catchment and enhanced accumulation of organic matter by litterfall from oak led to higher and more stable Hg accumulation rates under the oak-dominated forest.

Warm Period Caused Enhanced Mercury Uptake in Lakes Surrounded by a Mixed Forest (Section III.)

Core section III. represents the final stage of natural forest development in the Holocene, which was triggered by natural processes, namely, only by climate change, soil alteration, plant migration and competition [52]. At the second forest change, i.e., during the transition from an oak-dominated forest to a mountain mixed forest, the lake water balance likely flipped from oligotrophic to dystrophic. This water balance change could have occurred of increased mire formation and acidification in the lake catchments and is indicated by a decline and disappearance of *Isoetes* (quillworts) [68]. Therefore, the input of DOC into the poorly buffered aquatic system increased. Rydberg et al. [32] observed that Hg concentrations in sediments increased with fir and beech as a result of higher Hg:C ratios in the catchment soils, which was also observed in this work (Figure 3). However, because of a decreased input of OM due to the decreased litterfall in mixed forests, there is no indication of increasing Hg accumulation rates after the second forest change at

Herrenwiesersee. In contrast, the predicted Hg accumulation rates show a slight increase at Schurmsee and a moderate increase at Glaswaldsee up to 4600 yr BP. There is a more pronounced rapid increase in the accumulation of Hg between 4600-3800 yr BP, which is accompanied by a change towards decreasing C:N ratios leading to a positive effect of CP5 on Hg accumulation, likely due to the in-lake uptake of Hg caused by algae. Particularly at lake Schurmsee, the C:N ratios decline rapidly from 15.7 (4542 yr BP) to 12.5 (4490 yr BP). This abrupt change marks a warm period of rapid climate change [79], along with a higher production of autochthonous organic matter by algae and higher Hg accumulation rates in the sediment.

Furthermore, this core section at Schurmsee is characterized by multiple soil erosion events (Figure 3.A) and a sharp decline in C and Hg concentrations (Figure 2.A), which are particularly pronounced at 4700 and 3500 yr BP. These two events mark climatic transition zones from cold to warm and moderately humid, to again cold climatic conditions, which have been observed via diatom analysis in several sediment cores in northeastern Germany [80–82]. A study from the Swiss alpine lake Meidsee shows erosion events in comparable sediment layers indicated by increasing concentrations of Ti and Mn and decreasing concentrations of Hg and Pb [34]. Interestingly, the predicted Hg accumulation at Schurmsee (Figure 4.A) decreases until the end of core section III. (3000-2600 yr BP) during colder climatic conditions to values comparable to the mean Hg accumulation rates of core section II., which was dominated by oak and did not show any significant coniferous tree populations. Moreover, the effect of CP5 on Hg accumulation becomes negative (i.e., induced by increased C:N ratios) over the end of this section. This change was caused by increases in the relative importance of terrestrial OM erosion and leaching from the forest soil to the lake sediments. However, this decline in the predicted Hg accumulation rate at the end of the Bronze Age is not observed at Glaswaldsee.

Distinct Traces of Grazing and Agriculture in a Mixed Forest (Section IV.)

In the Neolithic period, the first farmers began the transformation of the landscape of the Black Forest lowlands into the cultural landscapes of today [83]. At Schurmsee and Glaswaldsee, distinct human impacts can be observed in the Late Bronze Age and Early Roman Iron Age, indicating that grazing started at the same time as cereal cultivation at Schurmsee, while grazing activities near Glaswaldsee developed ~2000 years before the first signals of grain cultivation, i.e., pollen of rye. Besides, the presence of ribwort and rye, there is another indicator of human disturbances. The Pb signal remains relatively low and steady until core section IV. Subsequently, the concentration of Pb peaks significantly at ~2400 yr BP at Schurmsee and at ~2000 yr BP at Glaswaldsee. For the first time, the Pb concentration trend follow the Hg concentration trend (Figure 2). Significant Pb signals during the period of the Roman Empire are also reported from different lake sediments and peat bogs [34,84–86]. Approximately 2100 yr BP, Roman Pb mining became the most important source of

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atmospheric Pb pollution in central Europe [86]. Mariet et al. [84] found, by means of Pb isotopic analyses of the sediments of Lac de Longemer, which is only 109 km to the northwest from Glaswaldsee on the opposite side of the Rhine Valley, no evidence for early local mining in the Vosges Mountains and linked the observed Pb isotopic signals at ~2000 yr BP to Roman origins. However, the first increase in Pb at Schurmsee started at 2800 yr BP (Figure 2.A) and unlike the trends in the aforementioned Pb studies, the Pb signal in this sediment core does not drop back to a pre-settlement background value. During an excavation at Neuenbürg, which is only 33 km away from lake Schurmsee (Figure 1), Grassmann et al. [67] discovered unusually well-preserved smelting complexes for iron and steel. Dated to the Late Hallstatt and Early Latène period (2600-2500 yr BP), these findings provide the earliest proof of iron manufacturing in this region. However, without an extended Pb isotopic analysis of the two sediment cores, the historical sources of Pb cannot be disentangled here.

More certain are the increases in Hg accumulation related to human activities in both sediment cores, while the proportional composition of tree species is relatively steady. The observed proxies for local settlement activities (e.g., rye for agriculture and Pb for mining activities) increase. Subsequently, the negative effect of the human impacts (CP2) reduced or even changed to a positive influence Hg accumulation. Compared to the previous warm period in core section III., contrasting effects of CP5 are observed, which may indicate a change in the source of the sedimented OM due to local agricultural activities. Because pollen analysis from eight small headwater lakes in the northern Black Forest showed significant differences in the timing of indicators for grazing and agriculture, we conclude that the lake sediments reflect the vegetation situation of their surroundings, up to a distance of no more than 1-2 km, without any major pollen inputs from greater distances [52]. Furthermore, it is notable that the PCR based prediction of Hg accumulation decreases at both lakes during a known pattern of regional settlement history, i.e., periods of major demographic crises (Figure 4). The temporary decline in anthropogenic factors during these crises led to lower predicted Hg accumulations rates. This result is consistent with the observed decline of human activities based on the pollen record (i.e., less agricultural plants), e.g., during the Migration period caused by the decay of the Roman Empire [39].

Pollution from Mining and Industrial Activities in a Mixed Forest (Section V.)

Mercury and Pb are emitted to the atmosphere from both anthropogenic and natural sources. However, anthropogenic emissions of these metals have greatly surpassed natural releases to the atmosphere over historical times [33,87,88]. With few interruptions, the upper core section documents a continuous increase in human influences in the northern Black Forest. The locations and history of at least six former mining deposits (silver, lead, and zinc) near the lakes are well

described (Figure 1). The mining activity, e.g., near the town of Freudenstadt, is noted for the first time in 1267 [89,90]. The local ore mining affected the chemical composition of the lake sediment cores noticeably. A striking feature is the rapid and simultaneous increases in the Hg and Pb concentrations at Glaswaldsee after the charcoal maximum at ~1100 yr BP, which are followed by a strong event of soil erosion due to deforestation caused by local mining (Figure 2.B). Breitenlechner et al. [91] found increasing values of charcoal in a peat deposit affiliated with a boom of silver mining ~1500 in western Austria. An increase in the Hg export from the lake catchments to the sediments during Medieval mining activities can be caused (i) by Hg release as a mining by-product of the excavated zinc-rich ores [92], (ii) due to the use of Hg in silver mining to amalgamate and concentrate the precious metal [93,94], (iii) by the release of bound Hg from the topsoil and burning of biomass after forest clearances [95–97] or through the local production of charcoal for mining [91,98].

Up to the High Medieval period, the predicted and observed Hg accumulation rates increase rapidly (Figure 4). Moreover, the effect of human influences (CP2) on Hg accumulation becomes positive in the last core section. Declines in the predicted Hg accumulation occurred during the crises of the Late Middle Ages, such as famines and plagues, especially the Black Death epidemic of 1347 to 1353 [99,100], and during the Little Ice Age, when unfavourable climatic conditions were followed by transnational political conflicts that significantly decimated the population in central Europe [101,102].

Furthermore, another development in human history is crucial for the accumulation of Hg in lake sediments. Overall, the industrial revolution beginning in the 1850s in Germany was characterized by investments in railway construction and heavy industry [103]. Exemplary, in 1846 only 1,518 steam engines were registered in the territory of the German Customs Union. This number had already increased to 8,695 by 1861 [104]. During this short period, coal as a primary energy source increased from 25% to 47% of the total nationwide energy consumption and peaked at 93% in 1942 [105,106]. This rapid increase in coal combustion and the related additional release of Hg to the atmosphere [107,108] are reflected by an increase in the Hg concentration (Figure 2.A) and observed Hg accumulation rates (Figure 3.A) in the sediment core of Schurmsee. Here, only the Schurmsee core contains signals of geochemical transitions during the industrial period. This observation raises the question why this additional atmospheric Hg (Hg_{AT}) input is not observed in the predicted Hg accumulation rates. The reason for this aberration may be a missing proxy for the additional input of Hg due to the combustion of fossil fuels after the beginning of the industrial revolution began. As mentioned above, the Hg concentration was not entered as a variable in the model. Thus, the PCR led to an underestimated value for the predicted Hg accumulation in the last five sediment layers.

5. Conclusions

Our results indicate that the deposition of Hg in two lake sediments of the Black Forest prior to ~3000 yr BP was caused by natural processes only, mainly by the quantity and composition of the forest vegetation. Consistent with our hypothesis, the evolution of conifers in the catchments validated the higher Hg concentrations in the lake sediments. However, there was no evidence for a substantial increase in Hg accumulation rates caused by the dominance of conifers. More than ~2000 yr after a mixed mountain forest was established, anthropogenic factors such as increased soil erosion after deforestation, grazing, agriculture activities, and emissions from mining and industry triggered changes in Hg accumulation stronger than natural changes in the forest composition alone. Our study demonstrates that a valid prediction of Hg accumulation rates can be obtained by a multiproxy stepwise regression analysis of lake sediment cores with well-defined pollen and geochemical records. Furthermore, this statistical approach offers to define and quantify controlling factors on Hg accumulation in forested areas over time.

Conflicts of Interest: The authors declare no conflicts of interest.

Autor Contributions: MS performed the chemical analyses, the interpretation of the results, the graphical presentation and wrote the paper. GT investigated the palynological record of Schurmsee. MPR assisted with the statistical analyses and interpretation. MR conceived of the sediment core sampling project in the Black Forest and investigated the palynological record of Glaswaldsee. HB conceived of the funding acquisition and led this research project.

Supplementary Materials:

Figure SM-1 and SM-2
Table SM-1 and SM-2

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